EFFECT OF ELEVATION AND RAINFALL ON THE PHYSICO-CHEMICAL PROPERTIES OF THE SOILS OF THE HIGHRANGES OF KERALA

ΒY

K. CHANDRASEKHARAN NAIR, B. Sc., B. Sc. (Ag.)

THESIS

Submitted in partial fulfilment of the requirements for the Degree of Master of Science in Agriculture (Agricultural Chemistry) of the University of Kerala

.

DIVISION OF AGRICULTURAL CHEMISTRY AGRICULTURAL COLLEGE AND RESEARCH INSTITUTE VELLAYANI TRIVANDRUM

CERTIFICATE

This is to certify that the thesis herewith submitted contains the results of bonafide research work carried out by Sri K. Chandrasekharan Nair under my supervision. No part of the work embodied in this thesis has been submitted earlier for the award of any dogree.

Y.A.

(P. KUMARA PILLAI) PRINCIPAL

2 some up

(M.H. KOSHY) Addl. Professor of Agrly Chemistry and Agricultural Chemist.

AGRICULTURAL COLLEGE & RESEARCH INSTITUTE, VELLAYANI : TRIVANDRUM.

AUGUST, 1969.

ACKNOLLEDGEMENTS

The author desires to place on record his acep felt appreciation and thanks to :

Dr. M.M. Kosby, Additional Professor of Chemistry and Agricultural Chemist for suggesting the problem and for the unstinted help and inspiring criticism at every stage of this investigation and in the preparation of this thosis;

Er N.S. Money, Professor of Chemistry for his valuable advice in the planning and execution of this work;

Prof. P. Kumara Illai, Principal, for providing necessary facilities for uncertaking this study;

br. R.S. Iyer, Sri P.R. Ramasubramanian and Smt. T. Pankajakshy Amma, Junior Professors of Agricultural Chemistry for the keen interest shown in persuing the thesis and for offering constructive criticism;

Jri K.N.K. Nair, Chief Scientific Officer, Kannan Deven Hills Produce Co. Ltd., for the immense help and guidance given in collecting the soil samples from High Ranges;

The members of the Staff and his collegues of the Division of the Agricultural Chemistry for much help rendered;

The Government of Kerala for ceputing him for Postgraduate course which enabled him to undertake this investigation.

K. CHAHLRAS, KHAHAM NAIR.

SINFINOS

IV	NOISUJONOD GUA YAAMMUS	**0	T 6
Λ	DISCARTION	• • •	68
ΛI	STUCTH	***	96
III	SUCHTER UNA STATASTAM		54
II	HEATEN OF LITCHATURE		Þ
I	NOITOMOATNI	***	τ
HOTAAND			nova

Sponfilled IIV

-11:0111-

F ***

INTRODUCTION

INTRODUCTION

The characteristics of a soil are determined for the most part by climatic influences such as precipitation, temperature, evaporation, humidity and wind velocity, by the nature of the vegetative cover, by the composition of the parent material from which the soil is derived and by the topography of the land.

The importance of climate and elevation on the development of soil characteristics cannot be over emphasized. Jenny (1941) by attributing functional concepts to soil forming processes evolved the following equation depicting the role of climate on the development of soil properties.

	ន	8	f (mT) o,r,p,t
Whero	s	5	Soil property
	m	2	loisture
	T	*	Temperature
	0	8	Organism
	r	2	reli m
	D	cz	parent matorial
	t		time

The climatic factors express themselves through the moisture and energy they contribute to an environment. The rainfall, which is a component of climate, imprints its effect on the soil through the cycle of water which enters the soil and which is used for the production of organic matter or for hyuration and hydrolysis or is released by the reverse process and removed through evaporation, transpiration and percolation.

Soveral works to study the influence of climate and clovation on the development of soil properties have been undertaken by Hirth (1926), Costin (1955), Dames (1955) and Muir (1965). Among the investigations on Indian soils, the work of Sen and Deb (1941, 1949), key Chauchuri (1943, 1957, 1965) whir (1967) and Gopalaswaay (1967) deserve special mention. Vijayachandran (1963) and Whanaraj (1966) have studied the effect of rainfall and clovation on the properties of soils in Kerala and Physore States.

The present investigation to study the influence of rainfall and elevation on soil properties is confined to the High Ranges of Kerala where the elevation ranges from 600 matres to 2100 materes and rainfall varies from 325 Cm. to 625 Cm. The main objectives of the investigations ares-

- to study the physical and chemical properties of the soils as influenced by rainfall and elevation;
- (2) to assess the extent of soll deterioration due to high rainfull and elevation.

The investigation is of special practical significance in the management of the soils of the High Ranges. The High Ranges has a highly significant place in the economy of Kerela as the dollar carning crops especially tea are grown in this area. The Kannan Jevan Hills produce Co. Ltd., alono cultivate about 45,000 acres of land under tea. Every year the plantation spends about Rs.70-80 lakhs on manuring alone and even if 40-80% of the soluble nutrients are leached below the root zone of the crop, the loss is tremendous and of sorious consequence. It is hoped that this study will threw some light on the leaching loss of nutrients in the solls of the High Ranges and that the inferences drawn will be applicable to the high-level plantations of kerals.

REVIEW OF LITERATURE

REVIEW OF LITERATORE

It was Dokuchaev (1899) who first considered soil as the result of the combined activity and reciprocal influence of parent material, plant and animal organisms, climate, age of land and topography. Joffe (1936) classified the factors involved in the process of soil formation into two, viz; active and passive. The passive factors included parent material, topography, and time and the active factors were comprised of the elements of biosphere, atmosphere and hydrosphere.

Vanziji (1934) reported that in north-cast Transvael, laterite soils were formed on any kind of geological formation.

In the laterite soils of Mauritius Graig and Helais (1934) observed that under increased rainfall red soils were formed by means of lateritic decomposition without a high leaching factor, whereas brownish yellow highly laterite soils were derived from lateritic weathering followed by a high leaching factor.

Vander Merve (1935) studied the soils of Transwaal and found that two different soil types had developed under conditions of heavy rainfall and high humidity.

From a soil survey of Puerto-Rico Roberts (1942) observed that there was no specific relationship letve a rainfall zones and soil groups.

In the soils of Lundelkhand Mukerjee and Agarvel (1945) observed that hot summers followed by neavy rainfall favoured weathering and consequent leaching coupled with the topography of the region gave rise to different types of soils.

Viswanath and Ukyl (1944) observed that geological aifferences might be obliterated under the influence of climate and that soils with similar proporties wight be covoring the most varied rock systems. They divided India longitudinally into four bread divisions based on Mayer's E.S. quotient.

Costin (1955) observed that in Australia under conditions of increased precipitation the climatic sequence of soils on basalt wore chernozens, chocolate soils and transitional chernozens.

Dames (1955) describing the soils of Java noted the occurrence of red laterite soils, brown laterite soils and humus trown solls durived from similar basic volcanic rocks under different rainfall conditions.

On the soils of Scotland Muir (1955) noticed very pronounced effect of rainfall. Hill peaks, peaky poduols 5

and peaty gleis occupied the high rainfall regions whereas brown forest soils and non-calcareous gleis were formed in areas where the rainfall was lower.

Bonahue (1958) stated that any soil profile is the result of both the direct, as woll as the indirect action of centuries of climatic forces.

Mohr (1913) found that in Java, the average temperature was largely a function of elevation above sea-level. He divides the island into four elimatic zones based on elevation. Describing the soils of these zones he concluded that laterite was formed in tropical low-lands with abundant rainfall and podsolization occurred in the tropical high mountain zones with high rainfall and middle laterite temperature. In the hill zones with lower temperature than the low lands transitional stages between these two were found.

In the Kumacan hills Kukerjee and Das (1940) found that three different soil types had developed, namely podzols, brown forest soils and red loams. They have given descriptions of the soil formations under the temporate climate observed in that hilly area.

Jenny (1941) stated that soil properties varied in the same direction with increasing altitude. 6

Retzer (1948) investigated the characteristics of soils developed from basalt under successive climatic vegetational zones of Jestern Colorado with elevation varying from 5000 to 10800 feet. He found many of the soil properties altering regularly with variation in altitude.

Costin and Hallsworth (1952) observed that soil properties do not vary in the same direction with increasing altitude. They noticed that in the monarc region of New South Wales, at lower extremes, soil formation was limited by insufficient deleture and in upper extremes by insufficient temperature.

Lames (1955), studying the soils of Java, noticed the formation of red laterites at elevations telow 300 metres brown laterites between 300-1000 metres and humus brown soils at altitudes above 1000 metres.

REFECT OF HAIN FALL ON SOLL PROPERTIES

I <u>Rainfall and physical properties</u>

a) <u>Colour</u>

Ritcher (1931), describing the soils of Hawaii, stated that the soils showed all gradations of colour from almost pure yellow through the different shades of brown and

red to grey and black in the different rainfall regions.

Durairaj (1951), in his study of the soil types of south India in relation to their colour and physical properties, observed that the orange hus of colour was closely associated with the composition of clay minerels.

(b) Physical properties

The physical properties of a soil are dependent principally upon its celloidal content.

Keen and Rackzkowski (1921) investigated the relationship between porcerace and clay content and found that these two were positively correlated. The correlation between specific gravity and clay content was found to be negative. Volume of expansion was found to be directly connected to the percentage of clay.

Merchand (1924) determined the specific gravity, porespace, water holding capacity and volume expansion of soils by the Keen Rackzskowski wethod and observed that cortain of the above properties could be roughly correlated with the clay content of the soil.

Coutts (1929) studied the volume expansion of Natel soils in relation to their clay content and obtained a positive correlation between the two.

Jonny and Leonard (1934) observed positive correlation between annual rainfall and clay content of soils of temperate regions.

Wilcox (1939) reported that the sand content of soils was negatively corrolated with the maximum water holding capacity.

Joachims and Kandiah (1947) in their study of the soils of Ceylon found that a high correlation existed between clay content and the unter-holding capacity.

Costin and Hallsworth (1952) observed that in the Alpine humas soils of South Jales, silt and clay were a maximum at an elevation of 5500-5000 feet and that these decreased significantly with increase, as well as decrease of altitude. This was attributed to the optimum temperature and moisture conditions prevailing in the sub-Alpine regions at altitudes of 5000-6000 feet.

Sinha <u>et al</u> (1957), in their study of the morphology of some red soils of Ranchi in relation to topography, concluded that down the slope the depth of profiles increased, the texture varied from sandy to clayey and the structure changed from single grain to granular or blocky.

Harradine and Jenny (1958) observed that soil texture became finer with increase in rainfall as a result of more

intense weathering but the relationship was not linear. High rainfall is less effective on texture than low rainfall, possibly on account of higher run off at high precipitations.

Unnikrishnan (1961) stated that above 60" of rain the physical properties of laterite soils wer, not influenced by variations in rainfall.

In a study of Milgiri soils Mahalingam (1962) found that clay content has no relationship to rainfall. He found a significant correlation between clay content and water-holding cepacity and a positive correlation between loss on ignition and clay content.

Dhanaraj (1966) found that moisture and water holding capacity did not show any significant correlation with rainfall or elevation but a close correlation existed between water holding capacity and pore space.

II Reinfell and Chemical properties

Crowther (1930) found that the silica-sesqui oxide ratios of the clay fraction were correlated negatively with ruinfall.

Craig and Halais (1934) found that the silicaelupina ratio was lowor and the iron oxide-alumina ratio

constant in the high rainfall laterite areas of Mauritius.

In a comparative study of several "Indian soils, Ukyl <u>at al</u> (1944) found that there was considerable decrease in the silica - alumina ratios with increase in the annual precepitation.

Ray Chaudhuri and Hian (1944) observed a comparitively low silica-alumina ratio and silica-sesquioxide ratio for the west coast soils formed under conditions of high rainfall.

Tsuneo Tamura (1956) found that the clay content of paxton soil clays increased abruptly at rainfalls ranging from 20-25".

In the row soils of Mysore State Puttaswamy Govda (1960) observed that under high Rainfell conditions leaching of calcium from soils lowered the pH and the silicate structures tenned to disintegrate. Consequently the silica so released tended to accumulate in the lower horizons.

Unnikrishnan (1961) concluded that alumina and total sesquicxides increased and the silica sesquioxide ratio decreased progressively with increasing rainfall.

Dhanaraj (1966) foună în high level laterites a high percentage of the sesqui-oxide.

III Roinfall and soil reaction

In a study of Maquilling soils Aristorenas (1937) found no relation between the $p^{\rm H}$ and dopth of soil.

In the case of the red soils of Mysore State, Puttaswamy Gouca (1960) indicated that under high rainfall conditions leaching of calcium from soils lowered the pH. Under conditions of low rainfall the case was reverse.

In his study of the soils of Madras State Wa Liar (1963) had indicated that increase in rainfall resulted in a progressive increase in the hydrogen ion concentration and prevalence of soil acidity. He obtained a significant negative correlation between $p^{\rm H}$ and rainfall.

Dhanaraj (1966) found a profound influence for reinfell on soil reaction. As rainfell increase there was a gradual decrease in the cation saturation. In the more humid zones leaching ensumed significant proportions resulting in progressive increase in hydrogen ion concentration and prevalence of soil acidity.

IV Rainfall and nutrient status

a) <u>Nitrogen</u>

Alway <u>et al</u> (1916) noticed that in the loose soils of Nebraska nitrogen content was a linear function of the rainfall.

Sievers and Holtz (1923) demonstrated that in the soils of Washington the total nitrogen content increased with increase in altitude and roinfall.

In sub-tropical timber soils Jonny (1930) found that nitrogen contant was not influenced by humidity factors. He found that it increased logarithmically with increase in N.J. quotient in grassland soils.

Hockonshith and Tucker (1933) found that the altrogen content of rocky mountain soils of Colarado increased with increases in elevation.

Dean (1937) Young that nitrogen content increased with increase in rainfoli.

Jenny (1941) found that the genoral trend of the nitrogen-depth curve was exponential. He also found that nitrogen generally penetrated coeper into the soil with increase in rainfall.

From a study of 43 cultivated soil profiles distributed all over India Sen <u>et al</u> (1946) found no correlation between climate and nitrogen content.

Ray Chaudhuri (1957) found that both nitrogen and organic matter contout increased with rainfall in the high altitude soils of India. In a study of the effect of climate on the nitrogen content and organic matter reserves of Indian soils Jonny and Ray Chaudhuri (1958) found that these factors increased with increase in rainfall.

Harradine and Jenny (1958) found that generally soils were richer in nitrogen under conditions of increased precipitation.

Mutkar and Ray Chaudhuri (1959) found that in black and alluvial soils nitrogen increased with rainfall and N.S. quotient.

Unnikrishnan (1961), Mahalingan (1962) and Vijayachandran (1963) observed that nitrogen content increased with rainfall in the soils of Madras and Kerala.

Dhanaraj (1966) noted a close relationship between total nitrogen and rainfall in Mysore soils.

(b) <u>Posphorus</u>

Wohltmann (1901) found that the surface soils of Germany were some times richer and some times poorer in phosphates than the corresponding sub-soils.

Hopkins (1910) found in the prairie soils of Illinois and Iowa, a high percentage of phosphorus in the surface than in the sub-soils. Remenn (1911) noticed that the distribution of phosphate in successive horizontal soil layers was not governed by any universal rule.

Aluay and Islam (1916) observed that the proportion of phosphate was gene, ally smaller in the first and second foot than in the lower layers.

In a study of 10 typical profiles of the major soil groups of Kerale Brito Muthunayagan and Koshy (1951) noted that the level of total phosphorus varied from 0.024 to 0.826%.

Aldrich and Euchanan (1954) observed that 60% of the South California soils examined by them had higher total phosphorus values in the first foot than in the second foot suggesting that rainfall as ociated with the downward movement of phosphorus was of minor importance in these soils.

Karim and Khan (1956), in a study of the vertical distribution of phosphorus in the soils of -ast Pakistan, found that it increased upto a depth of 7" and thereafter decreased sharply upto 35".

Walker and Adams (1958), in their study of the organic matter of soils formed on similar parent material but under e range of climatic conditions, found that as rainfell increased the phosphorus content of organic matter tended to be lower.

Kanvar and Greval (1959) found that the total reserves of native phosphorus are greater in the plains than in the hills where the soils are acidic.

Wild (1961), in a pedelogical study of phosphorus in 12 soils of Australia, found no correlation between loss of phosphorus and rainfall.

Nambiar (1963) in his study of the soil groups of Madras State observed that total phosphorus increased with rainfall significantly. Ukanaraj (1966) also made a similar observation with regard to Mysore soils.

(c) Potassium

Hilgard (1914) found that the relative migration and leaching of potassium in the soils was acfinitely more pronounced in hunid regions with heavy rainfall than in the arid regions of scanty rainfall.

Son <u>et al</u> (1949) revealed that lower values of potassium for Indian soils are only to be expected in view of the prevailing leaching conditions leading to laterisation in the humid tropics.

In a study of the vertical distribution of potassium in the soils of East Pakistan Karim and Khan (1956) found that this element decreased to a depth of 7" and thereafter increased progressively upto 35". Dhanaraj (1966) found no correlation between potassium content and rainfall in the soils of Mysore.

(a) <u>Calcium and magnesium</u>

In a study of the Nebraska loss soils Alway (1916) found a pronounced negative correlation between annual procipitation and hydrochloric acid soluble calcium, as well as hydrochloric acid insoluble calcium.

Russel and Engle (1925) observed that the depth of the line horizon increased with increasing rainfall in the soil horizons of the Central province of America.

In a study of certain soils of Eastern Uashington Vlasoff <u>et al</u> (1937) found that the percentage of calcium in the soil colloids decreased and that of magnesium increased with increasing depth of the profiles.

Das <u>of al</u> (1946) found a high percentage of magnesium in the black type of soils of India. This element tended to increase with depth in the profiles studied.

Joshi <u>et al</u> (1960) should that rainfall had a definite effect on the free lime content of the soil and the depth of lime accumulation zone in the soils of Kolhapur district.

(e) Organic matter and C/H ratio

Desausouro (1796) was of opinion that climate factors were responsible for the existence of different organic matter levels in soils.

Smolick(1926) found that the chemical composition of humus was affected by climate. The amount of humus, however, remained unaffected by the climatic factor.

Graig and Halais (1934) pointed out that the organic matter content increases and the C/N ratio became wider with increase in precipitation in Mauritius soils.

In their invostigations on the semiarid soils of Cape Province, losse and Worhill (1935) found a slight increase in C/N ratio with increase in rainfall.

In Havailan soils been (1937) found that organic earbon increased with increasing nitrogen, rainfall and elevation.

In their study of Chinsura and Chahjahangur soils of the humid régions Satyanarayana <u>at al</u> (1946) found that the carbon content was either steady or increased slightly within the first three feet and then decreased in the fourth foot. The C/N ratio of the brown soils in the humid zones increased in the second and third foot and marrowed down thereafter. In an investigation of the chocolate soils of New South Wales Mallsworth <u>et al</u> (1952) found that the organic enrbon content of the surface horizon increased with increasing rainfall.

birch and Friend (1956) studied the organic matter status of Bast African soils and found that the main factor governing the organic matter and nitrogen content was rainfall.

Jenny and Raychaudhuri (1953) in their study of the effect of climate on the organic matter reserves of Indian soils observed that organic carton increases with increase in rainfall.

Butkar and Ray Chauchuri (1959) noted that the organic carbon in black and alluvial soils increased with rainfall and N.J. quotient.

In a study of the C/N ratio of some Hawaiian soils Blombarg and Holmes (1959) found that organic carbon and nitrogen increased with rainfall.

In Alberta soils Synghel (1960) found that the carbon-nitrogen ratios were related to climate and that they increased progressively from the warm arid areas to the cooler more hunic regions.

In his study of Indian laterite solls Unnikrishnan (1961) found a significant positive correlation between rainfall and organic carbon under humld conditions. Mahalingaa (1962) got a high positive correlation between loss on ignition and organic matter percentage of the soil and significant positive correlation between the percentages of organic carbon and nitrogen.

In his study of the major soil groups of Madras State Namblar (1963) observed a close relationship between organic carbon and rainfall which he attributed to the direct effect of increased vegetation.

Vijayachandran (1963) found a significant positive correlation between rainfall and organic carbon in Kerala soils.

In the soils of Mysoro State Dhanaraj (1966) found wide carbon-nitrogen ratios which showed a tendency to widen with increase in rainfall.

Premenathan and Durairaj (1966), in a study of 36 surface alluvial soils, found that these was a close relationship between organic carbon and ignition loss.

EFFECT OF ELEVATION ON SOIL PROPERTIES

Jenny (1941) stated that soil properties varied in the same direction with increasing altitude. Rotzer (1948) found that many soil properties altered regularly with variations in altitude.

I Physical properties

Costin <u>at al</u> (1952) observed a close relationship between elevation and mechanical composition in the Alpine humus soils of New South Waled. Stone and gravel were found to be a maximum at altitudes above 6000'. Coarse sand was a manimum at altitudes between 5500' and 6000'. Silt and clay were found to be a maximum between 5500' and 6000' which tonded to decrease both with increase and decrease in the altitude.

II <u>Chemical properties</u>

Ray Chaudhuri and Mukerjee (1941) found that the silica content decreased significantly with increase in elevation in the laterite soils of India.

Ray Chaudhuri and Chakraborthy (1943) concluded that, so far as laterito soils were concerned, both annual rainCall and altitude above sea-level possessed cignificant negative correlation with silica-alumins ratio of the clay fraction.

In his study of Nilgiri soils Mahalingan (1962) indicated that the contents of iron oxide and aluminium oxide increased with elevation. Dhanaraj (1966) found that high level laterites contain a relatively higher percentage of sesquioxide.

III <u>Nutrient Status</u>

a) <u>Nitrogen</u>

Sievers and Holtz (1973) have deconstrated that the nitrogen content of the soil increased with increase in altitude and minfall.

Hocken Smith and Tucker (1933) studied the relationship between elevation and mitro, on content of grass land and forest soils in the rockey nountains of Colorado. They found that the mitrogen content of the soil increased with increase in elevation. Deam (1937) had also reported a similar relationship in Hawaiian soils.

Unnikrishnan (1961) and Mahalingam (1962) optained a significant positive correlation between elevation and nitrogen content for Indian soils.

b) Organic Matter and C/N Ratio

Jensan (1929) found that C/N ratio increased with the increase in soil temperature.

Mclean (1930) obtained a higher C/N ratio for the soils of the tropics as compared to those of the temperate region Dean (1937) found that the organic matter content of soils increased with increase in elevation. He also noted that elevation was inversity proportional to the mean annual temperature.

Retzer (1948) noticed an increase in organic matter content in the spils of Western Colorado with increaseing altitudes.

Costin <u>et al</u> (1952) found that the maximum organic matter content was obtained between 5000' and 6000' in the Alpine humus soils of South Wales. There was a fall in the organic matter content with decreasing elevation.

Jenny and Ray Chaudhuri (1958) observed a higher content of organic matter at higher elevations in their study of the effect of climate on the nitrogen and organic matter reserves of Indian Soils.

Unnikrishnan (1961) observed the inter-relationship between elevation and nitrogen in Indian laterite soils to be positive and significant.

Mahalingam (1962) found a positive correlation between elevation and the C/N ratio in the high level Nilgiri Solls.

In Mysore Soils, Bhanaraj (1966) found the relationship between organic carbon and elevation to be positive above an altitude of 500'.

MATERIALS AND METHODS

4

e

,

.

MATERIALS AND METHODS

The soils used in this study were collected from the High Hanges of Kerala situated in the Devikulas Taluk of Kottayan District. This is an interesting region both from the ecological and geological points of view. Though lying between 9° and 10° North latitudes, the average day temperature does not rise above 80°P even during mid-summer and between January and February the night temperature usually falls to 32°F with severe ground frosts. This unusual climatic conditions are mainly due to the elevation of the region and the characteristic configuration of the land which gives it a peculiar microclimate different from the surrounding place.

The main parent rock in this region is an archean igneous type rich in feldspars. Due to the rainfall ranging from 375 cm to 625 cm followed by three to four months of dry period there has been considerable laterisation and the resulting soil exhibits wide variations from argillaceous clay in the swamps to latosols of different colours, texture and composition in the uplands.

The total area of the High Ranges is 375 to 500 Sq.Km with a maximum width of 50 Km. from West to East and a length of 300 Km. from north to south. The rainfall drops significant from west to east. In Kalaar and Rajamalay the average rainfall is about 600 cm. per annum while 30 km. east of it the annual precipitation is less than 160 cm. showing a drop of 15 cm. for every Km. This variation in the rainfall is attributed to the peculiar topography of the land with numerous parallel Spurs across the main, north to south ridge of the Western Ghats.

Details of the soil samples collected

Soil profiles were taken from different elevations having varied rainfall. The places solected for taking profile pits were in the natural forests in the suburbs of the Kanan Devan Hills Produce Co. Ltd. Profiles were taken from elevations of 600, 900, 1200, 1500, 1800 and 2100 metres. Three profiles were taken from each elevation from spots separated by distances of 3-5 Km. Samples were collected from depths 0-16 Cm. 15-45 Cm. 45-75 Cm. and 75-105 Cm.

The general characteristics of the samples collected are given in Table I.

Analytical methods

The samples collected were dried in the shade, powdered gently and passed through a 2 um, sieve. The material under the sieve was stored in stoppered bottles for subsequent analysis.

Determination of physical properties

a) The mechanical composition of the soil was determined by the international pipette method given by Sankaram (1966).

b) Apparent density, absolute specific gravity, maximum water holding capacity and perespace of the solls were determined by the Keen naczkowski - box method as detailed by Sankaram (1966).

c) The pH of the soil was measured by using a Beckmann pH meter in a 1:2:6 soil water suspension.

Determination of chemical constitutents

a) Moistura

5 g. of the soil was taken in a weighing bottle and dried in an air oven at 105° C to constant weight. The difference in weight was noted and expressed as percentage on oven dry basis (Sankaran 1966).

b) Loss on ignition

10 g. of the soil was ignited at util red heat in a suffile furnace and the loss in weight determined and expressed as percentage on oven dry basis (Sankaram 1966).

c) Analysis of hydrochloric acid ortract

20 g. of the coll was digested with 6.75:3.25 hydrochloric acid on a sand bath for 2 hours and allowed to

settle over night. The extract was collected and made upto 500 ml. In the Hel extract the following estimations were made.

(1) Phosphorus

An aliquot of the Hel extract was evaported to dryness. 10 ml. of concentrated mitric acid was added and the contents evaported. The process of adding HSO₃ and Grying was repeated 3 or 4 times till the dried residue showed a granular appearance. The residue was extracted with dil Hel and phosphorus estimated in the filtrate by precipitating as ammonium phosphomolybdate as in the method described by Jackson (1958). The residue was expressed as acid soluble silica.

(11) Potassium

An aliquot of the acid extract was taken in a silica dish, evaporated to dryness and the residue ignited. The hot water extract of the residue was taken and K₂O estimated volumetrically by Volk's cobaltinitrite method as given by Piper (1950).

(111) Seen1-oxides

The sesqui-oxides were precipitated in 50 ml. of the acid extract using excess ammonium hydroxide. The residue was ignited and weighed and expressed as percentage as in the method given by Sankaram (1966).

(1v) Calcium and magnesium

The filtrate after separating the sesqui-oxides was made upto a known volume. In appropriate aliquotes calcium and magnesium were estimated by the versenate method as given by Mankaram (1966).

d) <u>Mitrogen</u>

This was ostimated by the Kjoldahl method (Fiper 1960).

e) Organic carton

This was estimated by walkley and black's method as given by Piper (1950).

Severation of clay and analysis of elay fraction

The clay was separated by the method detailed by Piper (1950)

The fusion extract of the clay was prepared by the method described by Sankaras (1966).

In the fusion extract the silica, sesqui exides, and aluminium exides were estimated by the methods given by Piper (1960).

Statistical Analysis

Simple correlations were worked out between different soil characoristics from the analytical data.

Table I

General Characteristics of the Soils

ŧ

Location	Profile and Lepth	Gravel	Colour	Texture
Anachal	<u>Profile I</u>		an an the state of	
600 motres	0-15 cm.	10.6	2.5 YR 3/6	Clay
	15-45 cm.	10.7	2.5 YR 4/8	Clay
	45-75 cm.	11.3	2.5 XI. 4/8	Clay
	75-106cm.	12.6	2.5 YR 4/6	Clay
	Profile II			
	0-15 cm.	10.9	7.5 YR 3/2	Clay
	15-45 cm.	3.4	5 YR 2/1	Clay
	45-75 cm.	3.6	7.5 YR 2/0	Clay
	76-105cm.	3.6	5 YR 3/2	Clay
	Profile III			
	0-15 cm.	3.6	70 Nr 3/1	Clay
	15-45 en.	1.2	7.5 YR 3/2	Clay
	45-75 cm.	7.0	7.5 YR 4/2	Clay
	75-105 cm .	14.6	7.5 YR 4/4	Clay

(Contd..)

F Location	Bopth and	Gravol 55	Colour		Texture
	Profile I				
Chithirapu- rem	0-15 ca.	9.6	7.5	318 3/2	Clay
900 metres	10-45 cn.	16.1	7.5	YR 4/4	Clay
	46 -7 5 cm.	7.6	5	YR 5/8	Clay
	75-105en.	9,3	5	YR 5/8	Clay
	Profile II				
	0-15 em.	7.8	5	YR 3/3	Clay
	15-45 cm.	14.6	5	YR 3/4	Clay
	45-75 cm.	14.5	5	A 4/8	Clay
	75-105en.	14.8	5	YR 4/8	Cley
	rofile Ill				
	0-15 em.	13.9	5	R 3/2	Clay
	15-45 cm.	8.4	5	YR 4/6	Clay
	46-75 cm.	5.7	5	<u>M. 4/8</u>	Clay
	76-105em.	4.8	6	YR 4/6	Clay

Table 1

General Characteristics of the Soils

(Contd..)

Table 1

General Characteristics of the Soils

Location	Profile and Depth	Cravel S	Colour	Texture
Kalasr	Profile I			
1200 metres	0-15 cm.	20.4	7.5 YR 7/2	Clay loan
	15-45 cm.	14.3	7.5 XR 8/2	loam
	45-75 cm.	14.6	7.5 m 8/2	loam
	75-105en.	18,3	7.5 YR 8/4	Silty loam
	Profile II			
	0-15 cm.	22.2	5 YR 4/6	Clay loam
	15-45 CR.	33.0	5 YR 6/6	Loan
	45-75 cm.	30.7	5 YR 6/6	loam
	75-105em.	17.3	5 YR 6/4	Silty cluy loan
	Profile III			
	0-15 ca.	61.9	7.5 YR 6/4	Clay loam
	15-45 cn.	14.9	5 Ya 6/8	loam
	45-75 ca.	24.3	5 YR 6/8	loam
	75-105cm.	10.8	S VR 7/4	oilty clay loam.

(Contd..)

General Characteristics of the Soils	General	Character	<u>listics</u>	of the	e Soils
--------------------------------------	---------	-----------	----------------	--------	---------

Location	Profile and Depth	Gravel	Colour	Texture
	<u>Profile I</u>			
Nullatan1	0+16 cm.	11.1	5 YR 3/3	Silty loam
1500 metros	15-45 cm.	10.6	5 XR 3/4	Silty loam
	45-75 cm.	16.9	5 YR 4/4	Silty loam
	75-105cm.	33,3	5 YR 4/4	Silty loam
	Profile II			
	0-15 cm	17.7	5 YR 4/6	Silty clay
	15-45 cm	17.7	5 YR 4/8	Silty clay
	45-75 cm.	12.8	5 YR 5/6	Cley loan
	75-105 cm.	7.9	5 YR 4/8	Clay loam
	Profile III			
	0-15 cm.	13.7	5 YR 3/4	dilty clay
	15-45 cm.	4.5	6 YR 3/2	Silty clay
	45-75 cm.	4.5	5 YR 2/2	Clay loan
	75-105 cm.	7.7	5 YR 3/2	Clay loam

(Contd..)

Table I

General Charact ritics of the Soil

Location	Profile and Depth	Gravel	Colour	Texture	
andre have a first and a first of the second se	<u>Profile I</u>	92.9945.09377.94579.09474.94674944762.0997			
Rajamalay					
1800 metres	C-15 cm.	0.0	5 xR 3/2	Clay	
	16-45 cm.	4.3	5 YR 3/4	Silty clay	
	45-75 cm.	13.4	5 IR 3/3	Clay	
	75-105 cm.	7.3	7.5 YR 4/4	Clay	
	Profile II				
	0-15 cm.	3.9	5 XR 3/2	Clay	
	15-45 cm.	2.0	5 YR 3/1	Clay	
	45-75 cm.	22.6	5 YR 3/2	Clay	
	75-105 cm.	11.3	2.6 YR 3/6	Clay	
	Profile III				
	0- 15 cm.	26.5	5 M 4/6	Clay	
	15-45 cm.	19.2	3.5 YR 3/6	Clay	
	46-75 ca.	15.0	2.5 YR 3/4	Silty clay	
	70-105 cm.	13.2	2.5 YR 3/4	Clay	

(Coutd..)

Table	ĩ
-------	---

<u>Goneral</u>	Charact	eristi	<u>cs of</u>	the	<u>Soils</u>

Location	Profile and Depth	Gravel Ş	Cold	our	Texture
	Profile I				
kejemalay					
2100 metres	0-15 cm.	7.9	5 1	R 2/2	Clay
	15-45 cm.	5.1	5 3	r 4/2	Clay
	45-75 cm.	20.1	7.5	R 4/4	Clay loan
	75-105 cm.	20.9	7.5	ar 5 /6	Clay
	Profile II				
	C-15 cm.	7.6	7.5	æ 4/4	Clay
	15-45 cm.	10.5	7.5 2	CR 4/4	Clay
	45-75 cm.	38.3	5 3	ER 4/4	Clay
	75-105 cm.	18.1	5 :	CR 4/8	Clay
1	rofile III				
	0-15 cm.	11.3	5 3	a 3/3	Clay
	15-45 cm.	21.1	5 3	CR 1/8	Clay
	45-75 cm.	11.6	5	(R 6/8	Clay
	75-105 cm.	30.4	5 1	CR 5/8	Clay

RESULTS

resolts

I. Physical properties

(1) Mechanical analysis

The results of the mechanical analysis of the soils are given in Tables II to VII and represented graphically in Fig. 1.

The highest percentage of coarse sand was found in soils collected from an elevation of 1200 metres (47.6%) and the lowest value (9.6%) was obtained in soils collected from an altitude of 1500 metres. The percentage of fine sand was highest in the soils developed at an elevation of 1200 metres (28.6%) and was lowest in the soils developed at an altitude of 1500 metres (2.2%). The silt fraction showed the maximum value in the soils at 1200 metres' elevation (34.4%) and the lowest was found in soils from 600 metres' elevation (3.0%). The clay content ranged from 14.7% in the soils at 1200 metres to 64.0% in the soils at 1600 metres.

The content of coarse sand increased with increase in depth at all elevations except in the soils formed at 1800

TALLE II

<

*

 \checkmark

Profile and D	epth (Cm.)	Coarse sand S	Pino Sand S	511t %	Clay %	Organic Batter
Profile I	0-15	27.8	10.7	3.2	52.4	3.9
	16-46	35.4	12.6	3.8	46.9	1.4
	45 -7 5	38.7	11.3	5.7	42.6	1.7
	75-105	37.8	13.0	5.5	42.5	2.2
Profile II	0-15	28.4	13.2	6.0	46.5	6.9
	15-45	33,3	14.9	4.3	40.5	7.0
	45-75	35.9	14.3	4.9	37.7	7.2
	75-105	37.5	14.1	6.3	35.2	6.9
P rofil o III	0-15	25.6	8.8	10.5	63.9	4.2
	16-45	30.6	7.6	11.1	47.3	3.5
	45 7 6	35.3	7.1	12.9	42.7	2.0
	75-105	35.7	8.6	12.9	41.8	1.0
Mean of all						
Profiles	0-15	27.3	9.9	6.6	50.6	5.0
	18-45	33.1	11.7	6.4	44.9	3.9
	45 -7 5	36.6	10.9	7.8	41.0	3,6
	75-105	37.0	11.5	8.2	39.8	3.4
Mean of all D	epths .					
Profile I		34.9	11.6	4.5	46.1	2,3
Profile II		33.8	14.1	5.4	39.7	7.0
Profile III		31.8	7.2	11.8	46.4	2.7

Mechanical Analysis of the soils collected from elevation 600 metres and reinfall 325 cm.

rofile and Dep	oth (Cm.)	Coarse sond F	Fine sand %	51 1t	Clay pé	Organic Matter
Profile I	0-15	24.9	15.9	4.5	52.0	3.7
	15-45	26.3	14.0	5.0	52.5	2.2
	46-75	27.6	13.3	5.7	51.6	1.8
	75-105	29.3	12.6	6.0	50.2	1.9
rofile II	0~15	26.5	13,6	6.7	47.3	5.9
	15-45	28.3	12.6	8.3	47.7	3,1
	45-75	25.5	7.9	8.9	54.5	3.2
	75-105	31,9	14.7	9.2	41.1	3.1
rof i le III	0-15	20.5	10.9	114	54.1	3.1
	16-45	21.8	9,6	12.1	55.3	1.2
	45 -7 5	22.6	8*9	12.8	54.7	1.1
	75-105	24.7	8.2	13.0	53.1	1.0
ean of all rofiles	0-15	23.9	13.4	7.5	50.8	4.2
	16-45	25.5	12.1	8.5	61.8	5.8
	46-75	26.2	10.0	9.1	53.6	2.0
	75-105	28.6	11.8	9.4	48.1	2.0
lean of all Der	the					
rofile I		27.0	13.9	5.3	51.3	2.4
rof i le II		28.1	12.2	8.3	47.6	3.8
rofile III		22.4	9.4	12.3	54.3	1.6

Mechanical Analysis of the soils collected from elevation 900 metres and rainfall 425 cm.

TABLE III

TABLE IV

Profile and Dep	th (Cn.)	Coarse sand	Fine sand \$	51lt %	Clay %	Organic matter
Profile I	0-15	41.7	21.5	10.5	25 .6	0.7
	15-45	47.0	28 .6	9.4	14.7	0.3
	45-75	43,1	22.9	16.5	17.3	0.2
	75.105	27.3	21.6	34.4	16.5	0.2
P rofile II	0-15	41.9	23.4	9.2	22.9	2.6
	15-45	47.5	23.6	10.3	17.6	1.0
	45-75	40.3	£0 .7	15.5	22.6	0.9
	75-106	22.9	18.5	20,3	30.5	0.8
Profile III	0-15	37.5	16.4	15.5	28.8	1.8
	15-45	45.0	20.5	14.4	19.6	0.5
	45-75	38.4	18.7	21.5	20.9	0.5
	75-105	22.0	17.8	30.2	29,5	0.5
<u>Kean of all</u>						
Profiles	0-15	40.4	20.4	11.7	28.8	1.7
	15.45	46.5	24.2	11.4	17.3	0.6
	45-75	40.6	20.8	17.8	20.2	0.5
	75-105	26.1	19.3	28 .3	25.5	C.5
lean of all Dept	the .					
Profile I		39.8	23.6	17.7	18.5	0.3
Profile II		38.1	21.5	13.8	23.4	1.3
Profile III		35.7	18.3	20.4	24.7	0.8

Mechanic 1 Analysis of the soils collected from olevation 1200 metres and rainfall 625 cm.

TABLE V

Mechanical Analysis of the soils collected from elevation 1500 metros and rainfall 535 Cm.

Profile and Dupt	h (Cm.)	Coarse sand \$	Fine sand	511t \$	Clay ×	Organic matter \$
Profile I	0-16	10.9	3,4	19.5	60.2	5,0
	15-45	11.5	2.2	18,4	64.0	3,9
	46-75	30.3	22.8	12.5	21.4	4,0
	75-105	30.0	19.1	15.0	32,0	3.9
Profile II	0-15	9.6	4.4	20.3	58.8	6.9
	15-45	10.2	4.1	19.2	61.8	4.7
	46-75	30.1	25.7	15.5	24.9	3.8
	75-108	25.3	30 .7	15.3	36.7	8.0
Profile III	0-15	15.1	8.5	18.2	64.5	3.7
	15-46	16.5	7.2	17,5	56.5	2,3
	45-75	30.2	16.9	16.1	35.8	5*0
	75-105	29,1	17.4	18.2	36.1	0.2
<u>Mean of all</u> profiles						
	0-16	11.9	5.4	19.3	57.8	5.2
	16-45	12.7	4.5	18.4	60.8	3.6
	45-76	33.2	21.8	14.4	27.4	3.2
	76-195	27.8	19.1	16.2	34.9	5*0
Meen of all Dept	he					
Profile I		22.9	11.9	16,3	44.4	4.2
Profile II		18.8	13.7	17.6	48.5	4.4
Profile III		22.5	12.5	17.2	45.7	2.1

TALLE VI

Mechanical Analysis of the soils colleged from elevation 1800 metres and rainfall 540 Cm.

Profile and Dep	th (Cm.)	Coarce sand %	Fine sand %	silt %	Clay	Organic matter
Profile I	0-15	27.9	5.9	16.5	45.6	5.1
	16-46	12,6	2.8	24.3	54.6	3,7
	45-75	21.9	4.3	16.2	55,1	2.5
	75-105	34.4	6.1	10.6	47.1	1.8
Frofile II	0-15	20.4	9,2	13.5	48.9	8.0
	15-45	17.6	7.6	18.1	5 3 .3	3,4
	45-75	29.5	6.9	19.5	50,0	3.1
	75-105	30.2	5.2	16.3	45.4	2.9
Profile III	0-15	22.3	20.4	14.7	49.7	2.9
	15-45	16.6	8.9	23.3	51.0	0.2
	45-75	18.9	10.3	21.0	49.5	0.3
	75-105	30.6	12.6	15.2	41.3	0.3
Mean of all						
<u>profilas</u>	0-15	23.5	8.5	14.6	48.1	5.3
	15-45	15.6	6.4	21,9	52.9	2.4
	45 -7 5	20.4	7.2	18.9	51.2	2.0
	75-106	31.7	7.9	14.0	44.6	1.7
Mean of all Dep	the					
Profile 1		24,2	4.8	16.6	50.6	3.3
Profile II		22.2	7.2	16.8	43.4	4.3
Profile III		22.1	10.5	18.5	47.9	0.9

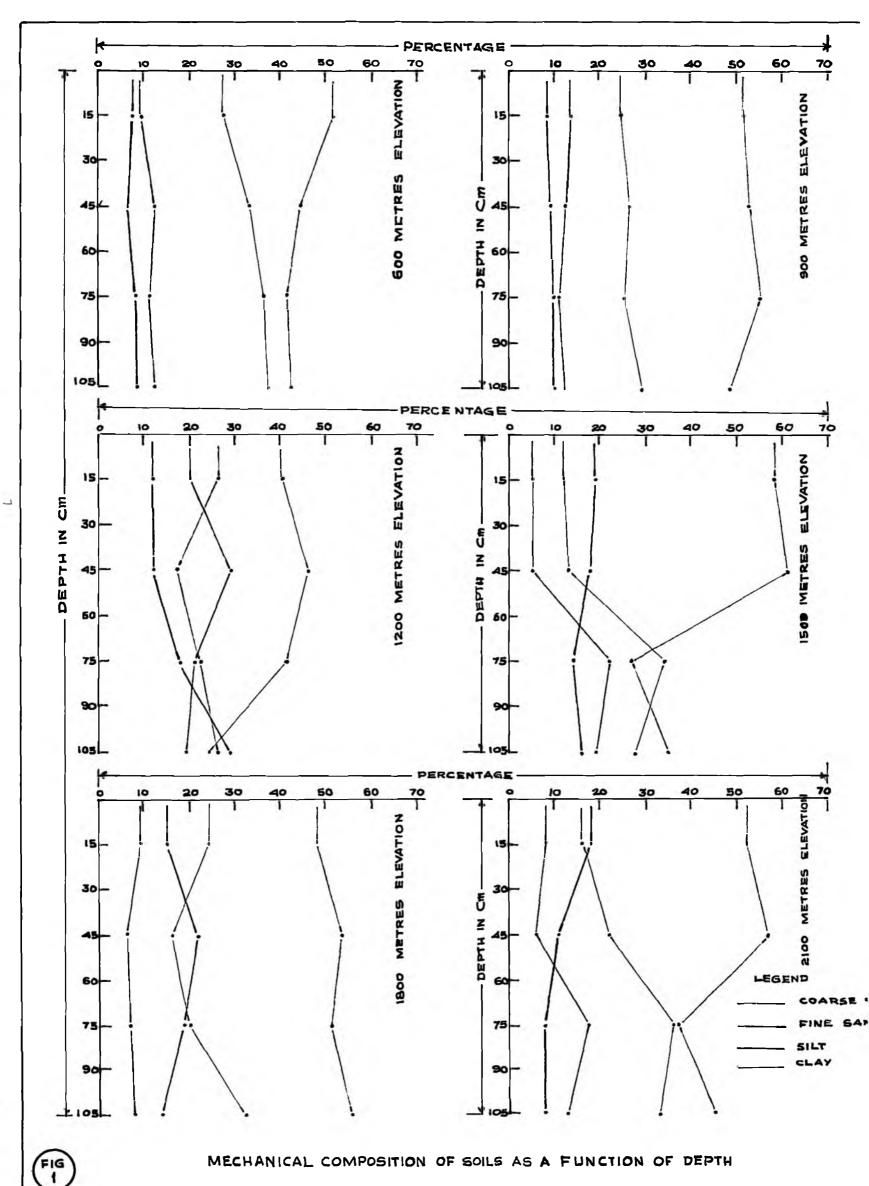
a la de la

(>

|--|

	sand Ş	Fine sond S	Silt	Clay S	Organic matter
0-15	12.2	6.3	18.4	58.2	4,9
15-45	P3.6	5.2	11.0	55.4	4.8
45-75	44.7	16.7	5.0	32.7	0.9
75-105	35.2	22.0	9.5	43.4	0.9
0-15	17.4	6.5	17.3	49.4	0.4
15-45	20.5	4.9	10.3	\$9.1	5.2
45-75	30.1	15.0	9.0	40.9	4.5
75-105	32.5	14.1	6.5	43.7	3.2
0-15	28.9	10.9	17.4	49.3	3.5
15-45	22.6	7.4	12.9	56.3	0.8
46-75	32.3	19.1	10.2	38.0	0.4
70-105	30.1	12.9	8.3	48.0	0.7
0-15	16.2	7.9	17.7	52.3	5.9
15-45	22.2	5,8	11.4	56.9	3.6
45-76	35.7	17.1	8.1	37.2	1.9
75-105	32.9	12.6	8.1	45.0	1.6
	28.9	9,8	10.9	47.4	5 9
	25.1	10.2	10.8	48.3	5.6
	25.9	12.6	12.2	47.9	1.3
	0 -1 5 16-45 45 - 76	0-15 16.2 16-45 22.2 45-76 35.7 75-105 32.9 28.9 25.1	0-15 16.2 7.9 16-45 22.2 5.8 45-76 35.7 17.1 75-105 32.9 12.6 28.9 9.8 25.1 10.2	0-15 16.2 7.9 17.7 16-45 22.2 5.8 11.4 45-76 35.7 17.1 8.1 75-105 32.9 12.6 8.1 28.9 9.8 10.9 25.1 10.2 10.8	0-15 16.2 7.9 17.7 52.3 16-45 22.2 5.8 11.4 56.9 45-75 35.7 17.1 8.1 37.2 75-105 32.9 12.6 8.1 45.0 28.9 9.8 10.9 47.4 25.1 10.2 10.8 48.3

Machanical Analysis of the soils collected from clovation 2100 metres and rainfall 550 cm.



1

-

-

TALLE VIII

Sincle value constants of from elevation 600 metres	the coils collected
from elevation 600 matres	and rainfall 325 cm.

?rofile	and	dep th(C m)	App arent donsi ty	Al solute specific gravity	Maximum water holding capacity	Pore- space
Profile	I	0-15	1.21	2,39	41.93	49.31
		15-45	1.41	2.64	33.01	46.59
		45-75	1.35	2,58	35.34	47.77
		75-105	1.41	2.59	30,98	43,87
Profile	11	0-15	1.25	2.43	42.16	40.74
		15-49	1.46	2.68	34.52	48.25
		45-75	1.38	2 ,63	36.14	49.12
		75-105	1.44	2,65	31.25	46.55
Profile	111	0-15	1.23	2.47	41.65	44.96
		15-45	1,32	2.69	32.14	45.24
		45-75	1.38	2.65	35.24	44 .26
		75-105	1,42	2.62	28.5 6	40.94
Yean of	<u>a11</u>					
profiles		0-15	1.13	2.43	41.91	48,00
		15-45	1.39	2.67	33,22	46.69
		45-75	1,36	2 .6 2	35.57	47.05
		75-105	1.42	2,62	30.26	43.45
<u>loan of</u>	<u>all</u>	Dopths				
Prof ile	I		1,35	2.65	35.32	46,88
Prof i le	11		1.38	2.60	36.02	48.16
Profile	171	1	1.33	2.61	34.39	43,85

TALLE IX

Single value constants of the soils of	ollected
from elevation 900 metres and rainfal	

Profile	and	Depth (Cm.)	Auparont density	Absolute specific gravity	Heximum vetor holding §	Pore- space
Profile	I	0-15	1.14	2,43	46.57	53.17
		15-46	1.24	2.54	41.20	41.84
		45-75	1.32	2.67	38.26	38,26
		75-105	1.33	2.70	38.49	39.14
Profile	II	0-15	1-12	2.40	45.46	52 .76
		1.6-45	1.22	3.62	40.29	40.25
		45-75	1.31	2 .6 1	37.14	37.52
		75-105	1.34	2,65	37.25	38.46
Profile	111	0-16	1.18	2.45	44.22	46.96
		26-45	1.24	2,56	39.54	40.24
		46-75	1.35	2.64	36.14	33.14
		75-105	1.37	2.72	36.56	40.25
Mean of	<u>a]]</u>					
pror 116:	ł.	0-15	1.15	2.43	45.41	60.96
		15-46	1.23	2.54	40.34	40.77
		45-75	1.33	2.64	37.18	38.31
		75-105	1.34	2 .6 9	37.43	39.28
Mean of	<u>all</u>	Depths :				
Profile	I		1.26	2.58	41.13	43.10
Profile	11		1.25	2.54	40.05	42,25
Profile	111		1.28	2.59	39.11	41.64

TABLE X

Single value constant of the soils collected from elevation 1330 metres as rainfall 625 cm.

					a series de la constant de la constant de la constant	
Profile	ana	Depth(Cn)	Apparent Density	Absolute specific gravity	Maximum water holding oapacity %	Pore- Space
Profile	I	0-15	1.25	2,51	37.10	49.13
		15-45	1.34	2.57	35 ,3 3	35.33
		45-75	1.28	2.48	30.55	38.55
		75-105	1.16	2.54	36.84	47.83
Profile	11	0-1 5	1.24	2.48	38.15	v0 . 25
		15-45	1.31	2.49	36.44	36.54
		45-75	1.25	2.47	31.12	39.72
		75-105	1.12	2.51	37.54	48.92
Prof i le	111	0-15	1.26	2.52	35.45	46.24
		15-45	1.34	2.56	33.92	36.25
		45-75	1.27	2,58	30.36	39.46
		75-105	1.24	2 .6 5	31.96	44.34
<u>Mean of</u>						
profiles		0-15	1.25	2.50	36.90	48.54
		15-48	1.33	2.54	35,23	37.04
		46-75	1,26	2.51	30,67	37.24
		7 5 10 5	1.17	2.57	35.44	46,99
Yeon of	<u>.</u>	Depths				
Profile	I		1.26	2.52	34.95	42 .71
Profile	II		1.23	2.49	35.81	43,85
Prof ile	113		1.28	2.68	32,92	41.55

TABLE KI

Single value constants of the soil collected from elevation 1900 metres and rainfoll 535 cm.

Profile and De	opth(Ca.)	Apparen t Density	Absolute specific gravity	Maximum vater holding capacity	Pore- space
Profile I	0-15	1.02	2,15	46,68	47.82
	15-45	1.09	2.17	45.08	46.30
	45-75	1.21	2.31	40,29	43,37
	75-105	1.32	2.38	32.19	43 .0 6
Profile II	0-16	1.02	2.18	45-68	48 .6 4
	15-45	1,11	2.20	44.04	46.82
	45~75	1.20	2,32	40.12	44.14
	76-105	1.35	2.36	30.18	42.18
Profile III	0-15	1.12	2,26	44.62	45.26
	15-45	1.15	2.28	43.82	44.04
	45 -7 5	1.18	2.32	40-12	42,57
	75.105	1.24	2,35	33.16	35.92
<u>Mean of all</u> profiles					
(particular) (C.107	0-15	1.06	2.19	45.66	47.24
	15-45	1.12	2.92	44.38	45 .6 9
	45-75	1,21	2.32	40.18	43.36
	76-105	1.30	2.34	31.84	40.39
Henn of all De	mths				
Profiles 1		1.16	2.24	41.06	46.11
Profiles II		1.18	2.26	40.05	45.44
Profiles III		1.17	2.30	40.43	41.94

TALLE XII

Single value constant of the soils collected from alovation 1800 metres any rainfall 640 cm.

P rofile s	anu	Depth (Ca.)	Apparent density	Al solute specific gravity	Maximum water holding capacity	Pore- space >\
Profile		0-15	1.14	2.19	42.29	45.49
		15-45	1.22	2.24	37.68	38,26
		45-75	1.35	2,28	32.63	36.15
		75-105	1.47	2,31	24.71	24.71
Profile	II	0-15	1.05	2.08	43.12	46,42
		15-45	1.18	2,21	38.24	40.29
		45-75	1.31	2,22	33.84	36,12
		75-105	1.38	2.24	25,24	26.15
Profile	III	0-16	1.02	2.18	40.25	42.56
		16-45	1.11	2.18	38,51	40.24
		45-75	1.18	2.28	31.24	33.96
		75-105	1.32	2.35	23.81	26.25
<u>Mean of a</u> profiles	11	0-15	1.07	2.14	41.88	44.82
		15-45	1.17	2.21	38.14	3 9 •5 9
		45-75	1.28	2.26	32.67	34.74
		75-105	1.39	2,30	24.59	25.70
Mean of a	122 1	Jepths				
Profile	I		1.29	2.26	34.33	35.90
Profile	II		1.23	2,19	35.11	37.24
Profile]	II		1.16	2.24	35.45	35.75

TABLE XIII

Single value constants of the soils collected from elevation 2100 metres and rainfall 550 cm.

Profile and	depth (Ca.)	Apparent density	Absolute specific gravity	Nexicus vator holding capacity	Pore- space
Profile I	0-15	1.25	2,12	40.26	41,12
	15-45	1.27	2.08	35,28	39.49
	45-75	1.29	2.05	28.31	28.31
	75-105	1.34	2,10	26.80	26,80
Profile II	0-15	1,82	2,04	42,25	45.12
	15-45	1.24	2.02	36.14	38.24
	45-75	1,25	2.01	29.26	30.54
	75-105	1,28	2.03	27.94	28,45
Profile III	0-15	1.28	2.12	39.61	42.94
	26-45	1,32	2.18	34.32	28 .66
	45-7 6	1.36	2.27	27.52	29.94
	76-105	1.38	2.35	24.21	26.46
Nean of all					
profiles	· 0 -1 5	1.25	2,09	40.67	43,06
	15-46	1.28	2.09	35.24	38,79
	45-75	1.29	2.11	28,36	29.59
	78-105	1.33	2.16	23 ,6 5	27.24
Mean of all	Denth				
Profile I		1,29	2.09	32.66	33.93
Profile II		1.25	2.03	33.89	30.69
Profile III		1.33	2.23	31.39	34.49

metres. At lower elevations the surface horizons generally contained a higher proportion of elay out at higher altitudes the clay content was more in the lower horizons.

(2) <u>Sincle value constants</u>

The single value constants of the soils are given in Tables VIII to XIII.

(a) Annaront dersity

The apparent density of the soils increased with depth in all the profiles except those collected from elevations of COO-1200 metres. The lowest value (1.02) was in the soils formed at 1500 metres and the highest value (1.47) was in the soils from 1600 metres' elevation.

(b) Alsolute appeific eravity

The absolute specific gravity increased with increase in depth in the soils at elevations from 900-2100 metres. The lowest value obtained was 2.02 in the soils from 2100 metres and highest value found was 2.72 in the soils at 900 metres' clevation. The absolute specific gravity generally increased with increase in elevation except at 600 metres and 1200 metres.

(c) <u>Maximum water-holding conectv</u>

The maximum vator-holding capacity of the soils varied from 24.7% in the soils at 1800 metres to 46.7% in the soils at 1400 metres. In the surface samples the maximum water holding capacity was 46.7% found in the soils at 1800 metres and the minimum value was 37.1% in the soils at 1800 metres.

(d) Pore space

The maximum pore space oltained was -3.2% in the soils formed at 200 metres and it decreased to 24.7% in the soils from 1800 metres.

3. Moisture

The moisture content of the samples is given in Tables XIV to XIX. It varied from 1.07% in the soils formed at 1200 metres to 13.05 in the soils at 2100 metres. At all elevations the moisture percentage showed a steady decrease with depth in all profiles.

4. Loss on ignition

The loss on ignition of the soils is given in the Tables XIV to AIX. The highest value, vi..., 35.44 was found in the soils at the elevation of 1800 metros. The soils at 1800 metros showed the lowest percentage of loss on ignition (3.8%). The loss on ignition proved a decrease down the profile at all elevations.

TABLE XIV

Profile and depth (Cn.)			pH	Noisture	Loss on Ignition
Profile	I	0-15	5.4	7.58	15.68
		15-45	5.4	4.13	13.54
		45-75	6.7	3.96	12.84
		75-105	5.5	3.62	12.04
Profile	II	0-15	5.6	6.42	14.62
		15-45	5.0	4.86	13.56
		45-76	4.5	4.12	11,93
		75-105	4.5	3,82	11.12
Profile	111	0-15	5.5	8.53	16.42
		15-45	5.4	5.26	15.25
		46-75	5.2	4.92	13.62
		75-105	5.0	4.16	12.96
Mean of	611				
proviles		0-15	5.5	7.51	15.57
		16-45	5.2	4.76	14.12
		45-75	5.1	4.33	12,76
		75.105	6 . 0	3.63	12.04
<u>Mean of</u>	all dop	the			
Profile	I		5.5	4.79	13.52
Profile	II		5.1	4.80	12.81
P r of i le	ĨII		6.3	5.71	14.53

pH. Mdsture and Loss on Ignition of the soils collected from elevation 600 metres and mainfall 325 cm.

TABLE	XV	

nH. Moisture and Lose on Ignition	1 of the soils
collected from elevation 900 metres	

Profile and De	opth (Ca.)	р ^Н	Moisture	Loss on Ignition
Frofile I	0-15	5.3	8.72	17.96
	15-45	5.5	8.10	15.06
	45-75	5.3	6.60	13.58
	75-105	5.6	6.00	10.26
Profile II	0-15	4.7	9.12	18.65
	15-45	5.0	8.71	15.56
	45-75	5.0	7.26	12.92
	75-105	4.9	6.82	11.02
Profile III	0-15	5,2	9.42	17.25
	15-46	5.1	8.86	14.96
	45-75	5.1	7.35	12.82
	75-105	5.1	6.72	10.89
<u>Nean of ell</u>				
profiles	0-15	5.0	9.08	17.95
	15-45	5.1	8.56	15.19
	46-75	5.1	7.07	13.11
	75-105	5,1	6,61	10 .7 2
Mean of all De	enthe			
Profile I		5.3	7.35	14.21
Profile II		4.9	7.98	14.54
Profile III		5.1	8.09	13.98

TABLE XVI

pH. Moisture and Loan on Ignition of the soils collected from elevation 1200 atras and mainfell 625 cm.

Profile	an a	Dopth	(Cm.)	p ^{ri}	Noisture ,3	Loss on Ignition %
l'rofile	I		0-18	5.4	3.98	13.20
			15-45	5.6	2.26	7.90
			48-75	5.5	2.02	4.16
			75-105	5.3	1.07	3.80
Prof ile	11		0-15	4.5	4.92	14.26
			15-45	4.5	2.56	9.62
			45-75	4.4	2,12	7.52
			75-106	4.4	1.62	4.16
Prof i lo	III	:	0-18	5.0	12.26	12.52
			15-45	4.9	10.12	7.85
			45-76	4.7	7.59	3.75
			75-1 05	4.5	6.42	2•93
<u>Mean of</u>						
profiles			0-15	4.9	7.06	13,32
			10-45	5.0	4.94	8.42
			46-75	4.9	3,92	5.14
			75-105	4.7	3.04	3.62
Mean of	<u>011</u>	denth				
Profilo	I			5.4	2.31	7.06
Profile	11			4.4	2.80	8,86
Profile	111			4.8	0.09	6.76

TALLE XVII

PH. Moisture and Less on Lemition of the Soils collected from elevation 1500 metres and rainfall 535 cm.

Profile and be	p th (C 2.)	в ^а	Noisture ø	Loss on Ignition
Profile I	0-15	5.2	11,28	30.70
	15-46	5.4	9,81	27.86
	45-76	6.4	8.75	22.06
	76-105	6.7	7.84	17.44
Profile II	0-15	4.7	12.52	29,73
	16-46	4,9	10.21	26,25
	45-75	5 .3	9.02	21.57
	75-105	5.4	7.53	16.24
Profile III	0-15	5.3	13.26	31.75
	16-46	5.4	11.46	28.24
	45-75	5.5	10.21	23.02
	78-105	5.6	8.54	17.62
Mean of all				19-10 MD-00
profiles	0-15	5.1	12,36	30.73
	15-45	5.2	10.49	27.45
	45-75	5.4	9.32	22.22
	76-105	5.5	7 .77	17.06
Mean of all De	otha			
Profile 1		5.4	9.27	24,21
Profile II		5.1	9,82	23,44
Profile III		5.2	10.86	25.13

<u>pH. Moisture and Lean on Lenition of the soils collected</u> <u>from elevation 1800 metres and rainfell 540 cn.</u>

Profiles and	Depth (Ca.)	рH	Moisture X	Loos on Ignition 3
Profile I	0-16	5.0	12.32	35,44
	18-45	5.0	8.93	23,66
	45 -7 5	5.5	8,23	13.90
	75-105	5.5	7.42	12.30
Profile II	0-15	5.4	13.53	32.94
	15-45	6.5	9.35	21.37
	45-75	5,7	8,96	10.24
	75-105	6.8	7.75	9.56
Profile III	0-15	5.2	12.24	33 .03
	25-45	5.3	9.16	23,94
	45-7 5	5.3	8.25	13.72
	75-105	5,4	7.16	11.95
<u>lern of all</u> profiles				
CEDETTER	0-15	5.2	12.99	33 .80
	15-45	5.3	9.14	22.99
	45-75	5.5	8.48	12.62
	75-105	5.6	7.44	11.27
Moan of all D	<u>epths</u>			
Profile I		5.2	9.22	21.32
Profile II		5.6	9.8 9	18.52
Profile III		5.3	9.20	20.66

TALLE XIX

pH. Moisture and Loss on Ignition of the soils collected from 2100 metres elevation and rainfall 550 cm.

Profiles and i	epth(Ca.)	рH	Mo istu re	Loss on Ignition %
Profile I	0-15	5.6	13.87	28.54
	15 -45	5.6	9.56	20.08
	4 5-7 5	5.8	8.73	19.63
	75-106	5.6	7.56	9.58
Prof il e II	0-15	5,3	12,93	26,54
	15-45	6,3	8.32	20.21
	46-76	6.4	7.25	19.66
	75-106	6.4	6,12	8,94
Profile III	0-15	5.2	13.52	27 .5 3
	15-45	5.4	9.24	20.26
	45-75	6 .6	8.12	18,64
	75-105	6.0	7.06	8.25
Mann of all				
profiles	0-16	5.3	13.44	27.63
	15-45	5.7	9.04	20,18
	45-75	5.9	8,03	19.26
	75-105	6.0	6.93	8.92
Moan of all De	<u>nthe</u>			
Profile I		5,6	9.93	16.47
Profile II		6.1	8.65	18.81
III efficient		0.5	9,48	24.86

There was significant negative correlation (r = -0.647) between loss on ignition and the sand contont and positive correlation (r = +0.541) between this property and clay content.

11. Chemical properties

The chemical properties of the soils are presented in Tables XIV to XXXVII.

(1) Soil reaction

The p^H of the soil (Tables XV to XIX) veried from 4.8 in the soils at 1200 metres to 6.4 in the coils at 2100 metres.

(11) Organic carbon

The organic carbon content of the soils (Tables XX to XXV) was found to vary from 0.14% at 1200 metres to 5.46% at 2100 metres. It decreased steadily with depth in all the profiles at all olevations.

The relationship letween organic carbon an. sand content of the soils was negative (r = -0.634) but a positive correlation existed between organic carbon and clay content (r = +0.382), as well as between organic carbon and water holding cpacity (r = +0.339).

TAPLE XX

Carbon, Mitrogen and C/N ratio of the soils collected from elevation 600 metres and rainfall 326 cm.

Profiles	and	dəpth (Ca.)	Carbon B	Nitrogen %	C/N Ratio
Profilo	I	0-15	1.268	0.114	11.1
		16-46	0.805	0.080	10.1
		45-75	0.965	0.065	14.8
		76-105	0.868	0,065	13.3
Profil le	II	0-15	4.010	0.262	15.3
		16-46	4.051	ú.174	23.2
		45-75	4.040	0.253	17.6
		75-305	4.921	0.239	17.5
Profile	111	0-16	2.472	0.142	17.4
		15-45	2.046	0.131	20.6
		46-75	1.08	0.095	12.8
		75-105	0.584	0,042	13.9
Nean of profiles					
	-	0-15	2.683	0,173	14.6
		15-45	2 .331	0.128	16,3
		45-75	2.071	0.1.29	16.1
		75-198	1.824	0.112	14.9
<u>Meon of</u>	<u>ell </u>	<u>)epths</u>			
Profilo	I		0.077	0.081	12.3
Profil le	II		4,030	0.223	18.4
Profi le	III		1.677	0.102	14.9

TALLE XXI

Profile and J	op th (Cm.)	Carton	Nitrojen X	C/N Ratio
Profilo 1	0-18	2.116	0+175	12.1
	15-45	1.239	0.166	7.5
	48-75	1.095	0.146	7.5
	75-106	1.130	0.182	8.2
Profilo II	0-16	3.362	0.304	16.4
	15-45	1.855	0.124	14.9
	45-75	1,853	0,121	15.3
	75-105	1.801	0.200	3 . 0
Profile III	0-15	1.773	0.121	14.6
	15-45	0.750	0.062	32.3
	45-75	0.640	0.060	10.8
	75-205	0.586	0.056	10.4
<u>Hean of all</u> p rofilos				
	0-15	2.417	0.166	14.36
	15-45	1.281	0.117	11.5
	45 7 6	1.196	0.109	11.2
	75-105	1.172	0.129	9.3
<u>Koon of All 14</u>	<u>opths</u>			
Frolite I		1.395	0.254	8.9
l'rofile li		2.218	0.162	13.9
Profile 111		0.937	0.074	11.9

Carton. Hitrogen and C/N ratio of the soils collected from elevation 900 metres and rainfall 425 cm.

TAULS RHIT

Corlon. Nitrogen and C/N ratio of the soils collected from olevation 1200 metros and Rainfoll 626 cm.

Profiles	ana	Dopth (C):::.)	Carl on	Nitrogen	C/N Ratic
Frofile	1	C)-15	0.431	0.086	5.0
		10	-45	0.171	0.044	4.0
		45	3 -75	3.137	0.028	4.9
		78	5 -105	0.137	0.030	6. 5
Profil o	11	()-16	1.514	0.131	11.5
		18	5-45	0.596	0.039	25.3
		49	5 -7 5	0.564	0.032	14.5
		79	5 -105	0.512	0.070	7.3
Profilo	111	c)-16	1.023	0.146	7.0
	1:	-46	0.259	0.032	8.1	
		4	5-75	0.278	0.035	7.9
		75	5-105	0.278	0.038	7.3
Moon of a profiles	977					
		C)-15	0.989	0.121	7.8
		14	8-45	0.342	0.038	0.1
		4	6-75	0.326	0.032	9.1
		7	5-105	0,309	0.046	6.3
Mean of	<u>.</u>	Denths				
Profile	ī			0.219	0.047	4.6
Profile	11			0.796	0.068	12,1
Prorile :	111			0,459	0.063	7.6

.

Carlon, N1(crog a and (G/N ratio of	the solls coll	ected from
el	ovation 1000) matres and	zninfa 11 536 c	m.

Profile	and	Dopth	(Ca.)	Carbon %	Hitrogen vi	C/II Ratio
ProSilo	I		0-25	8.937	0.709	4.1
			16-66	2.202	0.327	7.0
			45-75	2,350	0.440	5.2
			75 -1 05	2.278	0.205	11.1
Prof il e	31		0-18	4.025	0.257	16.6
		15-45	2,694	0.174	14,9	
		45-76	2,221	0.080	25.0	
		75-106	1.184	0.089	13.7	
. rofile	111	5	0-15	2.137	0.121	17.6
		15.45	2.363	0.175	12.1	
			45-76	1.128	0.102	10.9
			76-105	0.119	0.039	6.7
la n of						
profiles	ł		039	3.029	0.362	12,4
			10-45	2.126	0,004	11.3
			45-76	1.898	0.213	13.7
			75-106	1.193	0.104	10 .1
<u>Moan of</u>	<u>all</u>	denths	ł			
Profile	1			2.461	0.422	6,8
Profile	II			2.631	0.162	17.3
Profile	111			1.184	0.038	22.6

CONCUMPTION OF THE

anterina de la contra de la contracta de la con

Contracting and a state of the

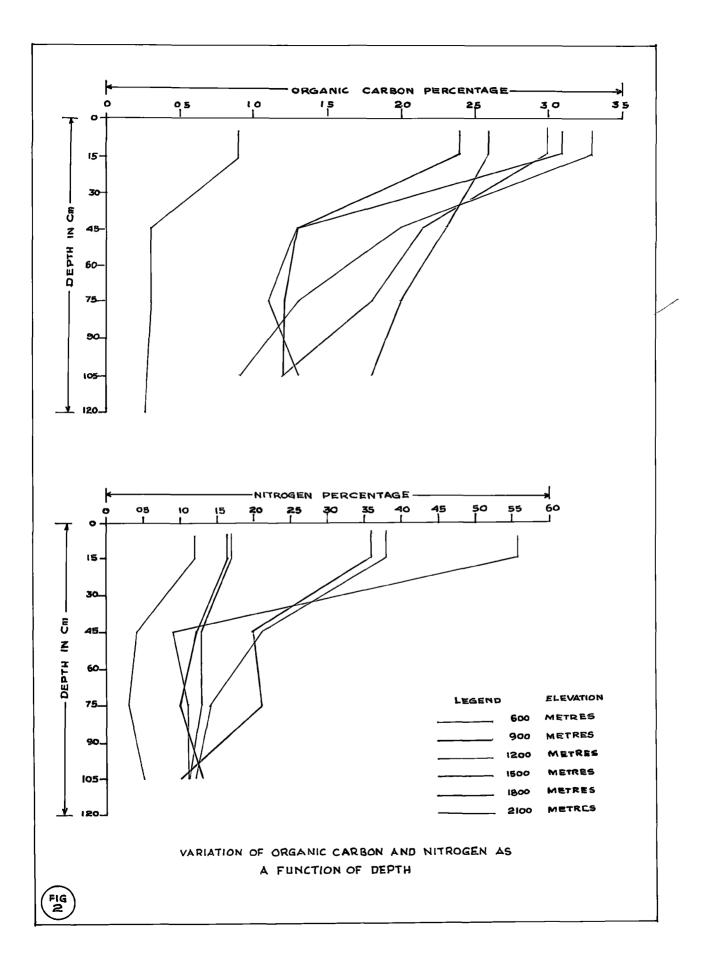
٩

<u>Carton. Nitrogen and C/N ratio of the soils collected from</u> elevation 1800 metres and rainfall 540 cm.

Profile and Dep	oth (Ca.)	Carkon Z	Nitrogen	C/N Ratio
Profile I	0-15	2.027	0.705	4.2
	15-45	3.105	0.164	12.8
	45-75	1.427	0.084	17.0
	7 5-105	1.051	0.092	11. 4
Profile II	0-13	4.682	0.851	5.5
	15-45	1,955	0.106	18.7
	45 -7 5	1.802	0.200	9.0
	75-105	1.684	0.210	8.0
Profile III	0-15	1.706	0.121	14.1
	15-45	0.132	0.028	4.7
	45-75	0.179	0.032	6.2
	75-105	0.199	0.035	5.7
<u>Mean of all</u> profiles				
	0-15	3.105	0.559	7.9
	15-45	1.397	0.099	12,1
	45 -7 5	1.142	0,105	7.4
	75-105	1.311	0.119	8,3
Mcan of all Der	oths			
Profile I		1.877	0.261	11.3
Profile II		2.531	0.341	10.3
Profile III		0.559	0.054	7.7

Carbon, Mitrogen and C/N ratio of the soils collected from elevation 2100 metres and rainbil 550 cm.

Profile and Dep	oth (Cm.)	Carbon	Nitrogen %	C/N Ratio
Profile	0-15	2,855	0.614	4.6
	15-48	2.764	0.346	7.9
	45-75	1.116	02167	6.7
	76-195	0.621	0.079	6,6
Profile II	0-15	5.402	0.380	14,4
	15-45	3.021	0.193	15.6
	45-75	2.031	0.191	13.7
	75-105	1.851	0.185	10.0
Profile III	0-15	2,046	0.141	14.5
	16-45	0.450	0.091	4.9
	45-75	0.245	0.061	4.8
	76-205	0.469	0.102	4.6
Mean of all				
<u>profiles</u>	0-15	3,454	0.378	11.2
	15-45	2.076	0.210	9.4
	45-75	1,331	0.136	8.4
	76-105	0.947	0.122	7.1
<u>Mean of all Da</u>	oths			
Profile I		1.609	0.301	6.4
Profile II		3.241	0.237	13.4
Profile III		0.802	0.096	7.2



In the soils there was significant positive correlation between elevation and organic carbon between elevations of 600 and 900 metres (r = +0.508), 900 and 1200 metres (r = +0.839), 1200 and 1500 metres (r = +0.581), 1500 and 1800 metres (r = +0.856) and 1800 and 2100 metres (r = +0.917).

(111) <u>Nitrogen</u>

1

The analytical results pertaining to the nitrogen content of the soils are given Tables to XX to XXV and graphically represented in Fig.2.

The nitrogen content showed remarkable variation with elevation and rainfall and ranged from 0.028% in the soils at 1200 metres to 0.709% at 1500 metres. This element decreased steadily with depth in most profiles.

The elevation - nitrogen relationship was positive in all the profiles at all elevations. The relationship between rainfall and this element was significantly positive in the surface layers, but negative in the deeper layers.

The organic carbon and nitrogen had positive correlation in the soils at all elevations (r = +0.385 at 600 metros, r = +0.553 at 900 metros, r = +0.826 at 1200 metros r = +0.926 at 1600 metros, and r = +0.339 at 2100 metros).

1:5

(iv) <u>Carbon - ditrogen ratio</u>

As seen in the Tables XX to XXV the C/N ratio varied from 4.0 in the soils at 1200 metres to as high a value as 25.0 at 1500 metres. The surface layer had the widest ratio when compared to the lower layers at all elevations.

(v) Acid insolubles

As shown in the Tables AXVI to XXXI the highest value for acid insolubles, vi..., 69.5%, was obtained in the soils at 1500 metros and the lowest value (44.4%) was found at the elevation of 1600 metros.

(v1) <u>Besoui-orioes</u>

The analytical data on the sesqui-oxide are given in Tables XXVI to XXAI. The sesqui-oxide content varied from 17.7% in the soils at 1800 metres to 42.3% in the soils at 600 metres. Vertical distribution of sesqui-oxides showed a stately increase down the profile in the soils at elevations of 600 metres, 1800 metres and 1800 metres but decreased with depth in profiles at elevations of 1800 metres and 2100 metres.

TABLE XXVI

Profile and D	opth (Cm.)	Acid insolutlos	Acid soluble silica	Sesqui- oxides
	1	· · · · · · · · · · · · · · · · · · ·	Ş	
Profile I	0-15	56.650	0.350	29.250
	15-45	51.260	0.410	32.625
	45-75	52,450	0.495	35.875
	75-105	46.600	0.160	39.150
Profile II	0-16		0.260	33.224
,	15-45	47.112	0.313	36.112
, ,	45-75	47.446	0.305	39.665
	75-105	42.112.	0.302	43,226
Profile III	0-15	53.115	0.215	32 ,584
	15-45	48-226	0.285	35.265
	45-75	49.109	0.346	38.275
	75-105	43.556	0.229	42,292
Mean of all p	rofiles		•	
	0-15	53,997	0.275	31,689
	15-45	48.866	0.336	34.667
÷	45-75	49.666	0.388	37.771
	75-105	44.089	0.230	41.556
Mean of all D	enthe			
Profile I		51.740	0.354	34.100
Profile II		47.224	0,295	38.057
Profile III		48.499	0.269	37.106

TABLE XXVII

Profile (and Depth	(Cm.)	Aoid insolubles %	Acid Soluble Silica	Sosqui- oxide Ø
			yo 		
Profile	I	0-15	51-130	0,460	27.150
		15-45	£0 - 560	0.575	31.500
		45-75	52.800	0.520	33.550
		75-108	51.225	0.550	36 .775
Profile I	1	0-15	48,120	0.275	30,112
		15-45	47.336	0,418	34,448
		45-75	49.230	0,423	37,665
		75-105	48.446	0.445	38.326
rofile 1	111	0-15	47.554	0,350	31.225
		15-45	46,226	0.495	34.115
		45-75	48,941	0,506	37,526
		76-105	47.212	0.521	38.229
ieen of s	11				
profiles		0-15	48.934	0.363	29.495
		15-45	48,041	0.496	33,354
		46-75	50.327	0.483	36.247
		75-105	48 . 9 6 1	0.505	37,410
lean of a	11 depth	8			
Profile	I		51,429	0.526	31.993
Profile	II		48,283	0.390	35.113
Profile	111		47.483	0.469	35.274

.

Acid incolutie and Acid solutie Silica and Sesoni-oxides in soils from elevation 900 metres and rainfall 425 cm.

Acid insolution on Acid solution Silice and Second-ovide in soils collected from elevation - 1700 metres and printell G25 cm.

Profile	ana nab t y	(Ca.)	Acid iasolyblos %	Acid Solublo Silica X	Cesqui- oxido %
Frofile	1	0-25	69.540	0.505	17,700
		15-45	63.120	0.430	18.770
		45-75	62.625	0.33	10.200
		75-105	60.990	0,340	23.400
Profilo	2I	0-15	64.536	0.150	22,766
		145	10.442	0.190	23.552
		43-75	60.486	0.223	24.226
		75-105	50.212	0.256	28.56.
Prolle	111	0-15	65.228	0.465	21,662
		15-15	6 9 . 929	0.431	22,782
		45-75	64.462	¢,286	23,880
		75-105	54.665	0.:24	27.662
lion of provide	<u>. 011</u>				
procle	10	0-15	66,442	0.373	20.706
		15-45	60.107	0.347	21,854
		40-76	63.624	0.281	22,435
		76-10 6	56.92-	0.273	26,642
<u>Sean of</u>	all unth	2			
Profile	1		64.826	0.402	19.767
Profile	11		6 64 9	0.204	04.874
Profilo	111		61,071	0.349	23.996

TABLO XXIX

Acid insoluble and Acid woluble Cilics and Second - Addes in the soils collected from elevation 1600 metres and eminfall 635 cm.

Profi le	anđ	Lepth (Cn.)	Acid Insolublas	Acid Solubles Silica	Sosqui- oxido #
Frofile	I	0-15	45+120	0.705	23,700
		15-45	44.350	0.465	28,600
		46-75	49.220	0.600	29,250
		75-105	49.780	0.530	30.800
Profile	11	0-16	49.0%	9.200	51 351
		15-45	48,902	0.095	26,996
		45-75	53.880	0.121	27.440
		7 5-1 06	63.006	0.140	22,226
Profile	113	0-15	41.652	0.495	27.685
		16-46	39.882	0.202	32,992
		45-76	40.112	0.116	38,656
		75- 205	42,226	0,132	33,862
Maan of	811				
profiles		0-35	45,299	0.460	24,195
		15-46	44,408	0.284	29,529
		45 -7 8	47.737	0.275	20.782
		75-205	48,667	0.267	30.962
Nean of	<u>all</u>	<u>konths</u>			
Profile	I		47.120	0.575	28.087
Prof118	11		61.623	0.139	26.971
Profile	111		40,043	0,258	31.793

TABLE XXX

Acid insoluble and Acid solublectlics and Sesoul-oxides in solls collected from elevation - 1800 metros and rainfall t 540 cm.

Frofile and De	epth (Ca.)	Acid Insolubles S	Acid Soluble Silico	Sesqu i- Oxidos S
Profile I	0-15	46.475	0.225	38,750
	15-45	51,180	0.110	36,562
	45-75	51.785	0.125	34,625
	75-105	50.730	0.275	33.660
Profile II	0-15	41.225	0.160	40.292
	15-45	49,080	0.088	38,992
	45-75	49,286	0.198	36.116
	75-105	48.992	0.208	35.552
Profile III	0-16	41.229	0.262	41.212
	15-45	48,990	0.145	39,112
	45-75	48.292	0.126	37,225
	75-105	47.926	0.172	36,112
Mean of all		10.000		
profiles	0-15	42,309	0.216	40.084
	15-45	49,750	0.114	38.222
	46 -75	49.788	0.149	35.988
	75-105	49,212	0.218	35,104
Mean of Depth				
Profile I		40.540	0.184	35.897
Profilo II		47.146	0,163	37.738
Profile III		46.609	0.176	38,415

•

·:9

rofile	ബപ	Depth	(Cm.)	Acid Insolubles	Acid Soluble Silica	Sesqui- oxides %
Profile	ï	nantið Handskarðsfr	0~15	60.560	0.233	32.250
			15-45	67.510	0.450	25,250
			45-75	58,125	0.350	24.760
			76-105	88 . 527)	0.300	24,500
?rofile	11		0-15	47.226	0.133	35,292
			15-45	64.116	0,208	28.662
			45-75	56,292	0.163	27.465
			75-105	55,612	0.150	26,662
.rofile	111	Č.	0-15	48.092	0.272	34.772
			15-45	65,225	0.205	28.115
			45-70	66.282	0.262	27.665
			75-105	59.116	0.245	25.012
<u>lean of</u>						
profile	8		0-16	48.926	0.213	36,438
			15-45	55.616	0.288	27,342
			45-76	56.566	0.258	26.626
<u>fean of</u>	<u>all</u>	denth	75 , 205	57.749	0.232	25.424
Profi 10	X			56,179	0,333	26.687
Profile				53.061	0.188	20.495
Profile	III			54,503	0.246	28,941

Acid insolutle and Acid soluble silica and Besaui-orides in Soils collected from elevation: 2100 metres and rainfall 550 cm.

TABLE XXXI

Elevation and sesqui-oxides were related significantly positive between 600 and 900 metres (r = +0.98), 900 and 1200 metres (r = +0.852) 1200 and 1500 metres (r = +0.766) and 1500 ' and 1800 metres (r = 0.442) but significantly negative at 1800 and 2100 metres (r = -0.652).

(vii) Phosphorus

The phosphorus status of the soils ranged from 0.091% at 600 metres to 0.293% at 2100 metres. The vertical distribution of phosphorus down the profile did not show any regular variation. The phosphorus content in the surface layers increased progressively in the soils from 600 to 900 metres and then decreased in the soils from 1200 to 2100 metres. The relationship between rainfall and phosphorus was positive at all depths at all elevations. This element showed no correlation with elevation at lower altitudes, but at higher altitudes, vi..., 1500 and 1800 metres, it showed a positive correlation (r = *0.571).

(viii) Potassium

The percentage of petassium as given in Tables XXAII to XXXVII, ranged from 0.061% in the soils at 600 metres to 0.51% in the soils at 1600 metres. There was a general tendency for the content of this element to increase with increase in

Profile	and Depti	a (Cm.)	P206 \$	K20 S	Ce0#	NgO 🖇
Profile	I	0-16	0.121	0.077	0.033	0.112
		15-45	0.104	0.090	0.017	0.108
		46-75	0.113	0.061	0.008	0.086
		75-105	0.091	0.188	0.011	0.055
Profile	II	0-15	0.116	0.072	0.096	0.259
		15-45	0.102	0.085	0.075	0.223
		45 -7 5	0.108	0.060	0,046	0.112
		75-105	0.030	0.182	0.035	0.076
Profile	111	0-15	0.125	0.085	0.062	0.185
		15-45	0.108	0,092	0.042	0.162
		45-75	0.115	0.064	0.035	0,108
		75-106	0.093	0.052	0.028	0.065
<u>Mean of</u>	all Prof	125				
		0-16	0.121	0.078	0,064	0.185
		15-45	0,105	0.089	0,045	0.164
		46-75	0.112	0.062	0.029	0.102
		75-108	0.091	0.141	0.025	0,065
<u>Kean of</u>	<u>denths</u>					
Prof ile	I		0.107	0.104	ć.017	0.090
Profile	11		0.104	0.099	0.041	0.823
Profile	111		0.1105	0.074	0.042	0.130

Prof. Kro. Cao Mgo in the soils collected from elevation 600 metres and rainfall 325 cm.

Th 147

TABLE XXXIII

P205, K20, Ca0, Mg0 in the soils collected from elevation 900 metres and rainfell 425 Cm.

P rofil os	and	Depth (Cm.)	P205\$	к20 %	Ca0%	MgO %
Profile	I	0-15	0.140	0.256	0.013	0.052
		15-45	0.150	0,256	0.011	0.092
		45-75	0.159	0.209	0.008	0.090
		76-195	0.156	0.208	0,008	0.089
Profile	II	0-15	0#120	0,254	0.053	0.069
		15-46	0.123	0,248	0.027	0.112
		45-75	0.125	0.001	0.013	0.112
		75-105	0,126	0.202	0.007	0.112
Profile	111	0-15	0.145	0.242	0.048	0.051
		15=45	0,158	0.241	0.021	0.092
		45-75	0.162	0,222	0.012	0.091
		75-1 05	0.161	0.212	0.007	0.082
<u>Mean of</u> profiles	11					
		0-15	0.135	0,251	0.038	0.057
		15-46	0,144	0.248	0.019	0.099
		45-75	0,146	0.211	0.011	0.098
		75-105	0.111	0.207	0.007	0.094
Mean of	lepti	19				
Profile	I		0.151	0,232	0.010	0.081
Profile	11		0.123	0.266	0.025	0.003
Profile :	III		0.156	0.229	0.022	0.079

,

P205, K20, CeO, HgO in the soils collected from elevation 1900 metres and rainfall 625 cm.

Profile	and Depth	(Cm.)	P205 8	к ₂ 0 %	Ca0 %	MgO S
Profile	ĩ	0-15	0.174	0,213	0.030	0,202
		16-45	0,178	0.280	0.022	0.102
		45-75	0.121	0.150	0.092	0.098
		76-105	0.167	0.240	0.022	0.098
Profile	II	0-18	0.152	0.202	0.060	0.266
		15-45	0.156	0.272	0.036	0.179
		48-75	0.115	0.142	0.051	0.262
		76-105	0.142	0.232	0.050	0.284
Profile	III	0-15	0.179	0.215	0.032	0.285
		15-45	0,182	0,225	0.048	0.162
		45-76	0.125	0.185	0.042	0.115
		76-105	0.171	0,242	0.042	0.099
Moan of	<u>all profi</u>	<u>l.es</u>				
		0-15	0.168	0.210	0.037	0.251
		16-45	0.172	0.269	0.035	0.147
		45-75	0.120	0.159	0.038	0.158
<u>Møsn of</u>	Denths	75-105	0.160	0.238	0.038	0.160
Profilo	I		0.160	0,221	0.024	0,125
P rofilo	II		0,141	0.212	0.046	0.248
Profi le	111		0.164	0.217	0.041	0.165

TABLE XXXV

P205,	К,0,	Cao,	MgO	in	the	60119	collected	from
and the second second	Carl Statements			in the second				

elevation 1500 metres and rainfall 535 Cm.

Profile and De	pth (Cm.)	P205 %	K20%	Ce0 🖇	MgO 🖇
Profile I	0-15	0.154	0.510	0.018	0.062
·	15-45	0,132	0.456	0.015	0.112
	45-75	0.140	0,463	0.013	0.092
	75-105	0.149	0.366	0.013	0.071
Profile II	0-15	0.135	0,412	0.025	0.077
	14-46	0.122	0.365	0.014	0.146
	45-75	0.131	0.372	0.030	0.085
	75-105	0.142	0.302	0.039	0.075
Profile III	0-15	0.148	0.380	0.022	0.065
	15-45	0.135	0.321	0.018	0.126
	45-78	0.126	0.312	0.014	0.116
	75-105	0.105	0.306	0.012	0,082
Mean of all pr		•			
	0-15	0.144	0.434	0.022	0.051
	15-45	0.129	0.381	0.012	0.128
	45-75	0.132	0.382	0.014	0.098
	75-105	0.135	0.324	0.021	0.076
<u>Mean of depths</u>					
Profile I		0.144	0.449	0.014	0.084
Profile II		0.132	0.363	0.027	0.096
Profile III		0.127	0.329	0.016	0.097

TABLE XXXVI

 P_2O_5 , K₂O, GoO, MgO in the soils collected from elevation 1800 metrics and rainfell 540 cm.

Profilo a	and	Depth (Cn.)	² 205 ^{با}	KgO ji	Ca) %	Mg0 \$
Profile	ĩ	0-15	0.167	0.443	0.060	0,062
		25-45	0.121	0.371	0.024	0.075
		46 -75	0.179	0.399	0.024	0.063
		75-105	0.176	0.435	0.016	0.065
Profi 1e	11	0-15	0.165	0.402	0.043	0.063
		10-45	0.108	0.331	0,040	0.080
		48-75	0.167	0.352	0.032	0.078
		75-105	0.162	0.392	0.029	0.077
Profile :	111	0-15	0.162	0.416	0.048	0.075
		15-45	0.142	0.386	0.034	0.092
		45-75	0,154	0.401	0.032	0.065
		75-205	0,135	0.376	0.033	0.065
Mean of	011	<u>profilos</u>				
		0-18	0.161	0.410	0.048	0.067
		16-45	0.123	0.362	0.033	0.082
		45-75	0.167	0.384	0.029	0.070
		75-105	0.157	0.401	0.026	0.069
Mean of	Deni	sha.				
Profile	ĩ		0.160	0,411	0.031	0.067
Profile	11		0.148	0.360	0.036	0.074
Piofilo :	111		0.148	0.394	0.036	0.074

TABLE XXXVII

Pg05, Kg0, Ca0, Mg0 in the soils collected from

٩

elevation 2100 metres and rainfall 550 cm.

-						
Profile (anu Dopt	:h (Ca.)	₽ ₂ 0 ₅ ≸	ä ₂ 0 \$	CaO 🖇	Иво 💈
Profi le	I	0-15	0.121	0.404	0.024	0.108
		16-45	0.293	0.359	0.016	0.143
		48-75	0,244	0.252	0,016	0.098
		75-105	0.199	0.305	0.016	0.086
Profile	II	0-15	0,112	0,359	0.046	0.119
		15-45	0.255	0.309	0,025	0,151
		45-75	0.212	0.302	0.038	0.106
		75-105	0.163	0.262	0.032	0.094
Profile	III	0 -15	0.225	0.365	0.018	0.108
		18-45	0.205	0.346	0.012	0.162
		45-75	0,212	0.026	0.010	0.104
		75 -105	0.202	0.340	0,008	0,098
Noon of	all prof	1105				
		0-15	0.152	0.376	0.029	0.112
		10-45	0.251	0.338	0.018	0 .16 2
		45-7 0	0.223	0.560	0.021	0.069
		75-105	0,184	0.302	0.019	0.092
Nean of	Venths					
Profile	I		0.214	0.0330	0.018	0,109
Profile	II		0.183	0.283	0.036	0.117
P rofile	111		0.211	0.344	0*015	0.118
			والمحافظ والمحافظ والمحافظ والمحافظ والمحافظ			

?7

,

elevation with the exception of 1200 and 1500 metres.

The correlation between rainfall and potassium content was positive at all depths of soil at all elevations ('r' ranged from +0.571 to +0.726).

(ix) <u>Calcium</u>

The calcium content of soils varied from 0.008% in the soils at 900 metres to 0.063% at 1800 metres. The calcium content decreased down the profiles at all the elevations. The relationship between calcium content and rainfall was positive at all horizons at all elevations ('r' ranges from +0.207 to +0.964).

(x) <u>Magnesium</u>

The magnesium content of soils is presented in Tables XXXII to XXXVII. The variation in magnesium content was from 0.051% in the soils at 900 metres elevation to 0.202% in the soils at 1200 metres. It tended to decrease with depth in most of the profiles.

The correlation between elevation and magnesium was negative in the soils at all elevations except at the elevation of 2100 metres where it was significantly positive (r = +0.84). The relationship letwoen rainfall and magnesium was negative at 0-15 cm and 15-45 cm, but was significantly positive at deeper layers. ('r' ranged from 0.549 to 0.607).

III Analysis of clay fraction

The characteristics of the clay fraction of the surface samples are presented in Table XXXVIII and represented graphically in Fig.3.

a) <u>8111ca</u>

The highest silica content was seen in the samples from an elevation of 2100 metros (48.55) and the lowest value (25.1,) was obtained at an elevation of 600 metres.

b) Alumina (Al_O_)

The highest content of aluming was seen in the soils at the elevation of 2100 metres (34.5%) and the lowest value of 22.9% was noted at the elevation of 1200 metres.

c) Iron oxide

The highest value of Fe_2O_3 (22.9,3) was found in the soils at an elevation 2100 metres and the lowest value (11.1,3) was obtained in the soils at an elevation of 1200 metres.

IV. Molecular Nation

The 310₂ : A1₂0₃ ratio varied from 1.72 to 2.39. The lowest value was obtained in the soils at 1200 metres elevation and the highest value was found at 2100 metres elevation.

 SiO_2 ; Fe_2O_3 ratio varied from 5.11 to 5.73. The soils at 600 metres' elevation has the lowest value and these at 1800 metres' elevation has the highest value.

SiO₂ ^R2^O3 Fattin varied at adily from 1.3 at 600 metres' elevation to 1.7 at 2100 metres' elevation.

PAPLE XXEVIII

١.

.

,

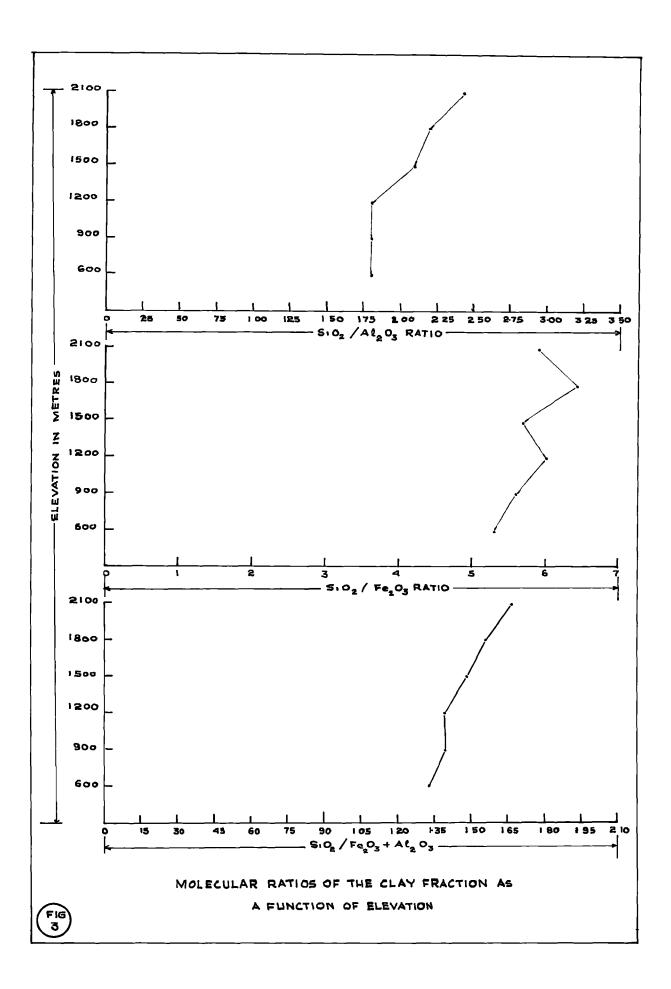
Chesical character of the clay fraction of surface soils

1

	١	Percentace on over dry basis				Molecular ratios		
Elovation	Profile	510 ₂	A1203	Fe ₂ 03	910 ₂	S102	\$10 ₂	
					1203	Fe203	Pe203 + 1203	
600 notres	I	29.71	28,25	15.22	1.79	5.21	1,33	
	II	30.41	29,15	15.93	1.77	5.11	1.31	
	III	28.51	27.51	13.12	1.76	5.79	1.35	
900 metres	I	27.22	25.82	13.42	I.77	5.42	1.34	
	II	25 .63	24.39	12.95	1.84	5.43	1.37	
	111	29.19	25.34	13.34	1.96	5.85	1.46	
1300 aetres	I	25.30	23,23	11.12	1.84	6.10	1.41	
	11	25.13	22.92	11.69	2.85	6 .91	1.41	
	111	27.26	26.83	12,23	1.72	5.97	1.34	
1500 motres	I	33.39	27.15	14.94	2.09	5.97	1.55	
	11	33.35	28.25	16.42	2.00	6.78	1.48	
	111	30.91	27.46	15.12	1.91	5.47	1.41	
1800 getres	I	37.84	29.21	26.13	2.21	6.24	1.63	
	II	33 .56	26.92	14.24	2.11	6.28	1.57	
	111	35.94	28.82	14.29	2.11	6.73	1.61	
2100 metres	I	48.50	34.45	22.94	2.39	S.65	1.68	
	11	44 .3 1	31.91	19.83	2.35	5.95	1.68	
	111	42.06	30.13	18.74	2.37	6.04	1.70	

<u>٦</u>

Ц



TALL AARIX

X

ACCORDED TO NO. 10

INCOME AND INCOME AND INCOME.

Physical and Chemical properties of the soils

Property (Average of all profiles and depths)	609 mitres	900 Rotres	1209 Motres	1500 metros	1800 Døtres	2100 Betres
I. <u>Mechanical Cornosition</u>						
(a) Coarse sand \$ (b) Fine sand \$ (c) Silt \$ (d) Elay \$ (e) Organic matter \$	33.6 11.0 7.3 44.1 4.0	25.9 11.8 8.6 51.1 2.6	38.1 21.3 17.4 22.4 0.8	21.4 12.7 17.1 41.2 3.6	22,9 7.6 17.4 40.3 2,8	28.6 10.9 11.3 47.5 3.3
II. <u>Sincle value constants</u>						
 (a) Appront density (b) Atsolute specific gravity (c) Maximum water holding capacity (d) Percentage of pore space \$\$ 	1.4 2.6 %35.2 53.0	1.3 2.6 40.1 42.3	1.3 2.5 34.6 42.7	1.2 2.3 40.5 44.2	1.2 2.2 34.3 36.3	1.2 2.1 32.6 34.7
II pH IV Noisture 5 V Loss on ignition 8 I Carton	5.3 5.1 13.6 2.19 0.14 15.6 0.107 0.092	5.1 7.6 10.9 1.52 0.13 11.7 0.143 0.242	4.6 4.7 7.6 0.43 0.00 8.1 0.155 0.216	5.2 9.9 24.4 2.06 0.22 9.3 0.134 0.350	5.4 9.4 20.2 1.66 0.22 7.5 0.152 0.331	5.7 9.4 21.0 1.95 0.21 9.3 0.202 0.219
I 420 S II 1120 J III S202/AL ₂ 03	0.933 0.147 1.77	0.019 0.087 1.86	0.037 0.176 1.80	0.029 0.092 2.00	0.034 0.071 2.14	0.021 0.114 2.37
IV S102/re203	5.37	5.56	5.96	5.74	6.41	8 .8 8
V \$102/Fe203 + 11203	1.33	1.32	÷ /	1.48	1.00	1.68

۰.

1

*

DISCUSSION

DISCUSSION

The present investigation comprises a study of the effect of elevation and rainfall on the chemical and physical properties of the soils of the High Ranges in Kerala State. The range of elevation covered in the study is from 600 metres to 2100 metres. The annual rainfall of the region varies from 325 cm. to 626 cm., the highest rainfall being at the elevation of 1200 metres.

The soils exhibit a wide range of colours such as datk grey, dark brown, dark reddish brown, yellowish red and dark reddish brown. The dark colour of the surface horizon changes to red in the sub soil. The brown colour of the surface horizon may be due to the presence of iron exides with accumulations of humus. The colour is lighter at the elevation of 1200 metres and the intensity of the colour increases with increase or decrease in elevation. The higher rainfall at 1200 metres might be responsible for the hydration of iron exides and the consequent lighter colour. Graig and Halais (1934) also observed similar conditions in Mauritius soils and Ritcher (1901) moted that such gradations of colour are functions of rainfall.

The general features of the soils reveal that they are of the lateritic type. The silica - alumina molecular ratio of the clay of the surface samples veries from 1.76 to 1.85 as the elevation varies from 600-900 metres and from 1.91 to 2.39 when the elvation changes from 900-2100 metres. The lower regions may therefore is considered as consisting of laterites and higher regions of non-laterites according to the definition of Martin and Doyne (1927). According to haychaudhuri and Sulaiman (1940), the silica - alumina ratio of Indian lateritos may is even higher than 2. In that case the soils in general has to is considered as laterites or as lateritic.

The machanical analysis data and textural classification of the soils indicate that the soils are either clay or clay loams. The first two layers (0-30 cm. and 30-46 cm.) contain the highest amounts of clay and there is a gradual decrease in the clay content down the profile. Similar observations were also made by Ricardo (1958). At 1200 metres the clay content shows a succen decrease in the intermediate layers and the decrease continues to the lover depths. The clay content of the surface sample is more in the profiles from elevations other than 1200 metres. Considering the other properties of the soil also there is every reason to infer that this elevation is a critical one with respect of the physico-chemical properties of soils. Costin at al (1952) found in the Alpine soils such a critical elevation above and bolow which there was approciable variation in the content of clay. The data for clay analysis in this study are

similar to those obtained by Rajagopal and Idnani (1963) for Hilgiri soils. There is an increase in the content of clay from 1800 metres, where the rainfall is a maximum to 600 metres where the rainfall is a minimum. This may be due to difference in the vertical translocation of clay brought about by the differences in rainfall. Similar observations were also made by Govindarajan and Datta Biswas (1968) in the soils of the Machkund basin and by Mahalingam (1962) in Nilgiri soils.

The water holding capacity is found to be highest in the soils developed at 1500 metres and having a clay content of 60.2%. The soils with the lowest clay content has also the lowest water holding capacity. It can therefore be assumed that the clay content of the soil is related to the water holding capacity which agrees with the findings of Mahalingam (1962) for Nilgiri soils. Apart from this apparent relation, there is no positive correlation between the clay content and moisture holding capacity which agrees with the view of Sen and Deb (1941) that moisture contents of heavy soils may not have any relation to the clay content.

In the soils studied the pore-space varies from 23.85 to 53.25 which agrees with the findings of Burger (1923) who noted that the pore space of forest soils varied from 28 to 695. A positive correlation was also noted between the elay content and pore space. This may be due to she large propertion of capillary pores present in the finer fraction.

The loss on ignition which is an approximate measure of the organic matter content is positiv by correlated to the clay content. The cooler climate in these regions might have favoured a higher production of Organic matter, which is foflected in the relatively higher values for loss on ignition, Mahalingam (1962) also found such relations in Milgiri soils.

The highest acidity (p^H 4.4) is noticed at an elevation of 1800 metros where the rainfall is a maximum. The scility can be attributed to the long and continued bloching of soils. Therefore it should be inferred that rainfall has a more pronounced effect than elevation on the development of acidity. Similar observations have also been made by Dhanaraj (1966) in Mysoro soils and Nambiar (1963) in Mauras soils. The cause of high acidity of the soils may also be attributed to the higher addition of organic matter and production of organic acids by its decomposition. Hasselman (1926) also observed that forest soils developed acidity if the trees occupying the land contained a lower amount of bases.

The soils studied are charactorised by relatively higher amounts of carbon and nitrogen. It has been found that elevation

has a significant positive correlation to the contents of nitrogen and organic carlon. This may be ous to the high ucisutre content of the soils which favours a vigorous plant growth and conservently higher organic matter production at higher elevations. Furthermore, as the elevation increases the climate becomes cooler which favours accumulation rather than humification of arcanic matter. This observation is in agroement with the findings of Jenny and Raychauchuri (1958), Unnikrishnan (1961), Mahalingan (1962), Maakiar (1963) and Vijaya Chandran (1963). The prosent invostigation royals that nitrogen content of the soils is high and has a positive correlation with rainfall and clevition which is in accordance with the findings of Alway (1916), Alevers and Holts (1923), Dean (1937). Ray Chaudhuri and Sen (1957). Unnikrishnan(1961). Mahalingan (1962). Ray Chaudhuri and Anjanoyalu (1965). The organic carbon content of the soil profile decreases with depth at all elevations which is avidently and to the vertical translocation of organic matter from the upper to the lower horizons by leaching. Similar observations were also made by Powers (1962) in forest soils. The soils developed at 1900 metres, which has the caximum rainfall, contain only low levels of organic carbon and nitrogen. As discussed earlier this elevation is a critical one with gradation in soil properties in both directions and the low content

of organic carbon and nitrogen can only be explained as que to the high leaching conditions prevailing at this elevation.

The analytical data show that the earbon-nitrogen ratios of the soil varies from 4 to 25. According to Satynarayana <u>et al</u> (1946) and Yadav and Pathak (1963) a similar range is observed in the uncultivated and forest soils of India. The high C/N ratio observed in some of the soils might be due to the reduced activity of micro-organisms under the acid conditions and low temperature prevailing at the higher elevations. Ray Chaudhuri and Anjaneyalu (1965) also got similar results in a study of the foot-hill soils of the Himalaya**3**.

The present study shows that the curface soils are some times richer and some time poorer in phosphorus, than the corresponding sub soils. This is in agreement with the findings of Malker and Adam (1959) that a decrease in total phosphorus with increased degree of leaching was due more to a decrease in bulk density associated with higher levels of organic matter than to the actual loss of phosphorus. The phosphorus content of the soils varies from 0.091% to 0.295% which agrees with the findings of Britomutunayagam and Koshy (1951) for the soils of Kerala and Yadav and Pathak (1963) for the forest soils of India. At 600 to 1500 metres, there is a positive correlation between the phosphorus content and loss on ignition which is a rough measure of the organic matter content. The relation between phosphorus and nitrogen is found to be positive at 600 metres* elevation and there is no correlation between the two at higher elevations. These differences can be attributed to the difference in the quantum of organic matter naturally added to the soil and the varied nature of its decomposition products at the different elevations. In Nilgiri soils Gopalaswamy <u>et al</u> (1962) also found similar correlations between organic carton and phosphorus at elevations alove 1800 metres and between nitrogen and phosphorus at altitudes alove 600 metres.

The data on sesqui-oxides reveal that they have been leached from the surface soil to the lower horizons in all profiles except those developed at 1600 metres and 2100 metres. The extent of leaching of the sesqui-oxides varies at the differences in the amount of complexing agents produced as a result of organic matter decomposition in the leaching medium. (Bloomfield, 1966; Thomas, 1964).

The study shows that calcium content is highest in the surface soils and it decreases down the profile at all elevations. \This observation Cagrees with the findings of Bost and Zetterberg (1932) and of Thomas (1964). The high content of calcium in

à

the surface soil may be due to the accusulation of this element through leaf fall. Elevation has no appreciable influence on the content of calcium except at elevations above 1800 metres which may be due to the differences in the type of vegetation in the crea.

The analytical data reveal that magnosium content has no relation to elevation from 600 metres to 1600 metres, but above that elevation there is positive correlation between the two. The relation between rainfall and the content of magnesium is negative which may be due to the leaching conditions consequent on a high rainfall. The leaching of magnesium by rainfall is a perious problem affecting the productivity of the soils of this region.

The observations have revealed that 1200 metros is a critical elevation from where the Physico-chemical properties of souls change in both directions. It is also seen that the nature and content of organic matter of forest soils control to a great extent the distribution transformation and translocation of other mutrient elements.

SUMMARY AND CONCLUSIONS

SUMMARY AND CONCLUSIONS

A study has been undertaken of the soils of the High kanges of Kerala to investigate the influence of elevation and rainfall on the physical and chemical properties of the soils. The elevations from which the soils were collected varied from 600 metres to 2100 metres and the rainfall from 328 cm. to 625 cm.

The chief findings are summerised below -

- (1) The soils exhibit a wide range of colours such as dark grey, dark brown, reddish brown, yellowish red and dark reddish brown. The colour of the soil is lighter at 1200 metres due to the hydration of the sesquioxides consequent on the high rainfell received at this elevation.
- (2) The surface horizon generally contains the maxium anount of elay which decreases with depth down the profile. The translocation of clay is maximum in the soils collected from an elevation of 1200 metres where the highest amount of rainfall is received. The clay content of the soils is correlated to their water holeing capacity.
- (3) The highest actualty is noticed in the soils developed at an elevation of 1900 metres.

(4) All the soils are characterised by relatively higher amounts of carbon and nitrogen. There is significant positive correlation between elevation and the contents of these elements. The organic carbon contont of the soils decrease with depth at all elevations which may be due to the effect of leaching. The carbon-nitrogen ratio varies from 4 to 25.

- (5) The surface soils generally show a higher content of phosphorus than the corresponding sub soils. There is positive correlation between phosphorus content and organic matter in the soils from elevations of 600-1600 metres. The relationship between phosphorus and nitrogen is found to be positivo only at 600 metres.
- (6) The sesqui-oxides have been leached downward from the surface soils at all elevations except in the soils at 1800 metres and 2100 metres.
- (7) The calcium content is highest in the surface soils and it decreases down the profile at all elevations.
- (8) The magnesium content of the soils has no relation to elevation from 600 metres to 1600 metres, but there is a positive correlation above that elevation. The magnesiu content of the soils is inversely proportional to the rainfall.

- (9) The clay fraction analysis has shown that soils are laterities or lateritie.
- (10) The elevation of 1900 metres which has the highest rainfull is a critical one as far as soil characteristics are concerned because most soil properties are found to be either a maximum or minimum at this clovation.

REFERENCES

/____

REPURENCES

< ~

Alhisvar Sen, Ukyl, A.C. Mukherjee, S.K. and Visvanath, B	1946	Comparative study on Indian Soils IV Mechanical analysis of soil properties Ind. J. Aarig. Sol. 161 246-250
Aldrich, v.C. and Euchanan, J.R.	1054	Phosphorus content of soils and the parent rock in Southern California. Soil. Sci. 77 : 369-376
Alvoy, J.P.	1916	The loess soils of Nebrasia portion of the transition I, II <u>Soil. Sci.</u> 1: 197-208.
Alway, C.J. and Dlish, K.J.	1916	Loess soils of the Nabraska portion of the transitin region, humus, humus nitrogon and colour. Soil, scl. 1: 230-258
Alway F.T. and Islam, R.M	1916	The loess soils of Nebraska portion of the transition region IV Mechanical Composition and inogranic constitutents. Soil. Sci. 11 405-435.
Aristorenas	1937	A study of certain Physical and chomical characteristics of some Maguilling soils <u>Phillipp</u> , <u>Acric</u> , <u>26</u> : 542-559
Birch, H.F. and Priond, M.T.	1956	The organic matter and nitrogen status of East African solls. J. <u>Soil. Sci. 71</u> 156-167.

	Bloomberg, H.F. and Holmes, W.	1 959 '	The Carbon nitrogen ratio in some Hewaiian solls. <u>Hawaii. arric. Expt. Ste.</u> <u>Progr. Ren. 1. 221</u> 7
	Bloomfield, C	1955	Leaf leachates as a factor in Pedogenesis <u>J. Sci.Fd.</u> <u>Agric. 6</u> : 641-651
ŗ	Brito Mutunayagam, A.F.A and Koshy, M.M.	1951	Chemicel nature and distribution of phosphorus in soils Travancore-Cochin. Bull. Cent. Res. Inst. Univ. Travancore 21 68-76
	Burger Hans.	1923	Physikalische Eigensehatften der vald-und Freilandhoden Buhler-Buchdruck.Zürich pp.221.
• ,	Costin, A.B. Halls worth, E.G. and Noof, M.	1952	Studios on pedegeneds in New South Wales III Alpine Sumus soils, <u>J. Soil Sci.3</u> , 190.
	Costing, A.B.	1955	A note on the basalt soils in Britain and Australia <u>J. Soil</u> Soi. 6: 268-269.
	Coutts, J.R.H.	1929	Single value soil properties, a study of the significance of certain soil constants II. Studies in Natel soils. J. agric. Sci. 19: 325-341.
	Craig, N. and Halais, P.	1934	The influence of maturity and rainfall on the properties of laterite soils in Mauritius. Emp.J. Expt. Agric. 21349-359

Crowther, S.Y.	1930	The relationship of climate and goological factors to the composition of the soil clay and distribution of soil types. <u>Proc. Rov. Soc. Lull. 107</u> : 1-30.
Dames, T.V.G.	1955	The soils of Central Java. <u>Contributions of the General</u> <u>Agricultural Research Station</u> <u>BORDE</u>
Das,S. Kukhorjee, S.K. Abhiswar Sen and Viowath, B.	1946	Comparative Studies on Indian Solis III. Ease exchange proper- ties. Indian. J. Arric. Sol. 16 : 234-239.
Dean, L.A.	1937	The effect of rainfall on the Carbon and Nitrogen Contont any C/N ratios in Hawaiian Soils. <u>Soil, Joi, Foc. Amer.</u> <u>FFOST 2.</u> : 455-450.
Dhanaraj, J.	1966	A study of the mutrient status of Mysore State soils as influenced by election and rainfall. <u>Dissertation</u> <u>Submitted to and accepted by the</u> <u>University of Madras</u> .
Dhir, R.P.	1967	Pedological characteristics of some soils of North Vestern Himalayas. J. Indian. 200. Soil. Science. 13.: 61-69.
Dockuchaev, V.V.	1839	Quoted by Jacob.S. Joffe in A.L.C. of soils.

Durairaj, D.J.	1051	Study of the soil types of South India in relation to their colour and physical and chemical properties. <u>Thesis submitted to the</u> <u>University of Madras.</u>
Ghani, N.9. and Alcen, S.A.	1943	Studies on the distribution of different forms of phosphoru in some Indian Soils II Jepth distribution. Ind. L. Arric. Sci. 13:377-381.
Gopalaswany, A. Natarajan, R. Rengasuany, P. and Gopalakrishnan, V.	1968	Relationship between loss on ignition, Mitrogen and Phosphor in the Alluvial and later its soils of Madras State. J. Ind. Soc. Soil Sci. 16: 379-382.
Govinderajen, S.V. and Datta biswas, H.R.	1968	Characterisation of certain soils in the sut-tropical Hunid Zones in the South Sustern part of India-soils of Machkund Easin. J. Ind. Soc. Soil. Sci. 16:2: 170-186.
Nallsworth, R.G. Costing, Al., Collins, F. R. and Robertson, G.K.	1952	Studies on podogenesis in Nev South Vales <u>J. Soil Sci. 3</u> : 193-124.
Harradino, F. and Jonny, H.	1958	Influence of parent material and climate on texture, nitrogon and carbon contents of Virgin California Soils.L. Texture and nitrogen contents of soils.

Sot1. 3e1. 85 : 235-243.

/

* Hosselman Henrick	1926	Studios over Larrskogens humistacke, dessegenskapor och Leroendi an skogsverden Middol. f. <u>Statene Skogsforsek-</u> anstalk. 22 + 169-552.
Hockonsmith, B.D. and Tuckor, E.	1933	The relation of elevation to nitrogen content of grass land and forest soils in hocky neuratins of Colorado Soil, joil 26: 41-450
Hopkins C.G.	1910	<u>Soil fertility and permanent</u> <u>Acriculturo</u> Ginn and Company Loston.
Issec, W.E. and Gorhill, B	1936	The organic matter content and Carbon nitrogen ration of Some semiarid soils of Cape province. Frans. Nov. Acc. Jouth Africa. 231 245-254.
Jackson, M.L.	2958	<u>The Soil Chewical Analysis.</u> MeGrav Hill Book Co. Inc. New York.
Jonny, X.	1928	Relation of climatic factors to the amount of nitrogen in soils. J. Amer. Soc. Agron: 21. 900-912.
Jenny, H.	1930	The Nitrogon Content of the solls as related to the precipitation evoparation ratio. <u>Soll. Sol.</u> 22: 193-207.

Δ.

Jenny, H. and Leonard.	1934	Functional relationship between soil properties and rainfall soil. Acl. 281 363-381.
Jonny, H.	1941	Factors of soil formation <u>Mogray Hill Co. New York and</u> London.
Jonny "H.	1950	Causes of high nitrogen and organic matter content of certain tropical forest soils. <u>Soil. 301.631</u> 2-6.
Jonny, H. and Ray Chaudhar1, S.P.	1958	Affect of climate and cultivation on the organic matter poserves of Indian soils <u>ICAN New Bolbi.</u>
Jonsen, H.L.	1920	On the influence of Carbon nitrogen ratios of organic material on the mineralisation of nitrogen. <u>J. Azric. Sci. 19.</u> 71-75.
Joachim, A.W.R. and Kandiah, S.	1947	Studies on Ceylon soils XVII and XVIII. The physical and physico-chemical characterstics of the major soil types of Ceylon. <u>Tron. Astic. 103</u> : 71-84.
Joffe, J.S.	1949	A.B.C. Soils.
Joshy, K.V. Pharende, K.S. and Dixit, L.A.	1960	Spils of Sugar Cane areas in Kolhapur District. J. Ind. Soc. Soil. Sci. 9. 157-167.

Kanwar, J.S. and 1959 Forms of phosphorus in Punjel Growal, J.S. Soils. Ind. 306. Soil. Sci. 71 135-146. 1056 Vertical distribution of Karim, A. and Nitrogen, phosphorus and potassium in some soils of Khan, D.H. last Pakiston. Soll. Sel. 81: 1-6. Keen, L.A. and 1921 The relationship between clay Reckzkowski. cont at any c.rtain physical properties of soils. J. Acric. Sci. 11: 441-449. Mahalingan, Pok. 1961 A study of the physical and Chemical properties of the high lov-1 Nilgiri soils Dissertation for Diff. (As.) Subsitted in and accrited by the University of Madras. liartin, F.J. and Doyne, H.C. Latorite and lateritic solls 1927 of Sierra Loone. J. AETIG. Sci. 17: 533-547. 1930 The carbon mitrogen rutio of Nelean, W. Soil organic matter. J. Agric. Sel. 20: 348-356. 1924 Some physical properties of Merchand, L.D. Pransvaal Soils. J. Agric. Sci. 14: 151-169.

vii

Mohr. ...C.J. 1913 weoted by senstius M.V. Agro geolo ical studies in Tropics.l. hi, h altitudes oriental tropics. Soil. Res. 2: 10-54. Huir, Jaw. 1965 The offect of soil forming factors over an area in the South of Scotlana. J. Soll. Sci. 6: 84-93. Mukherjee, L.K. and Studies on Bundelkhand Soils I. The genetic 1943 Agarwal. R.R. types. Ind. J. Acric. Sci. 13: 587-597. Mutkar, V.K. and Carton and Nitrogen status 1969 Ray Chaudhuri, S.P. of arid and Semiarid regions of India. J. Ind. Soc. Soil Sci. 71 255-262. Paul I. Vlasoff and Characters of certain soil 1937 Lawrence. C. Meeting. profiles of South Eastern Jashington. Goll, Mel. 14: 65-81. Piper, J.S. 1950 Soil and plant Analysis The University of Adelaide Adelaide, Austraila. Characterstics of forest soils 1932 Povers, ...L. of North Jestern United States. Soil, Sci. 34: 1-10. Premenathan, S. A study of the Alluvial 1966 Durairal, D. soils of Haarus state. Macres. Leri. J. (2): 49-54

- Puttaswamy Gowda, E.S. 1960 Investigations on Red Soil formation under different climatic conditions and parent material. Dissertation for the avard of Postgraduate Diplona in L.A.R. I., New Delhi. 1 Rajagopal. C.K. 1963 Some aspects of phosphorus Idnani, H.A. fortilization in the Nilgiri soil. J. Ind. Boc. Soil. Sci. 21 142-150. Ramann 1911 Quoted by Das, S. in Vertical distribution of phosphate in calcareous soils. Ind. J. Agric. Sci. 16: 164-168. Raychaudhuri, S.P. and Sulaiman. M. 1940 Studies on the Chemical constituents of Indian laterite anu red soils I Determinatin of free sesquioxid components. ind. J. Agri. Sci: 10 (2) 158-163. Haychauchuri, S.P. Studies on Indian red soils 1941 III General morphological and lukheriee. H.R. characteristics of some
- Raychaudhuri, S.F. and Chakraborthy, J.N. 1943 Studies on Indian red soils IV influence of rainfall and altitude above scalevel on the Chamical composition of the clay fraction of the soil types. Indi. J. Asric. noi. 13: 252.

profiles.

Ind. J. Apri. Sol. 11: 236-240.

- Naychaudhurl, S.P. and Mian, N. 1944 Studies on Indian rod soile VIII. Studies on the physicochemical and minerological properties of some Indian red and laterite soils. Indian L. Arric. Sci. 141 117-129.
- Raychauchuri, S.P. 1267 As reported by Rahoja in <u>Soll Productivity and grop</u> <u><u>arowth</u>.</u>
- Haychaudhuri, J.F. Anjaneyalu, E.S.R. and ShuMa, S.S. 11L-122.
- Retzer, IL. 1948 Soils developed from Basalt in Western Colerado. <u>Koil. Sci. 66</u>: 365-375.
- Ricardo, E. Jahn 1958 The Guataparo Red colls of Northern Venezulla. Soil. Sei. 56: 1-6.

*Ritcher, G.

*Roberts, R.C.

- 1931 Physical properties of Eavali Solls with special reference to the Colloidal fraction. <u>Havalis Asrics Expts Sts Fulls</u> 621 1-45.
- 1942 Soils of puerto Rico U.S.D.A. <u>hur. plant. Indust: Ser.</u> 1936 No.2.

- Roy. L. Donahue 1958 <u>Soils, an introduction to</u> Soils and mlant grouth. Engle wood Cliffs, N.J. Prontice-Hall INC. Rost, C.O. 1932 Replaceable bases in the Jean, M and Soils of Eastern minnesote and
- Schlarborg.
- Russel, J.C. and 1925 Joil Horizons in the Central Engle, L.J. Prairie. Rost. Fifth Meeting <u>Amer. Soil. Survey. Assoc.</u> <u>61</u>1-18.
- Russel, J.C. McRuer, Am.C. 1927 The relation of organic matter and nitrogon content to Series and types in Virgin grass land soils.

Soil. Soi.: 24: 421-451.

the effect of lime upon them.

Boil. Sci. 33:

- Sadanandan Nambiar. 1963 A study of the wajor soil groups of Madras State. <u>Dissertation submitted to and</u> accepted by the University of Madras.
- Smolick 1926 Quoted by Janmy, H. <u>Holl. Sol. 201</u> 1950.
- Sankaram, A. 1966 <u>A latoratory Hanual for</u> <u>Agricultural Chemistry</u>. Asian publishing House, Dontay.
- Batyanarayana, K.V.S. 1946 Swaminathan, K. and Viswanath, B. Comparative studies on Indian Soils carbon and nitrogen status of Indian soils and their profiles.

Ind. J. Agrio. Sci. 16: 1946.

Sen. A.T. Ashutosh. and Laterite and Rod Scils of Bhupendra Chandra Dob 1941 India II. Ind. J. Acri. Sol. XI: 617. Sen, A.T. Potassium Utatus and availability 1949 Del. b.C. and to Crops of non exchangeable Bose, S.K. potassium in some Indian Red and latorite soils. Soil Sci. 68: 291-293. Shewan, J.M. 1938 Quoted by Russel. Soil conditions and plant growth . Longmans Green and Co. Ltd. 1923 Siovers. J.T. and Holtz, H.F. The influence of precipitation on soil composition and soil organic matter maintenince. Mash. Agric. Sxat. Sta. Bull. 1761 1-30. Sinha, 3.D. A study of the morphology 1.957 Vorma, h.r. of some of the red soils of Lall, A.L. Ranchi District in relation to the topography of the land. J. Ind. Soc. Soll. Sel. 6: 956-972. Synghol, K. M. 1960 Charactoristics of fifty Alberta soils. J. Ind. Soc. Soil. Sci. 8: 77-83. 1964 Studies on some forest soil of Thomas, K.M. Kerala. Thesis submitted to and accepted by the University of Kerala. Tsungo Tamura 1956 Physical, Chemical and minorological properties of Brown podzolic soils in Southorn New England. Soil. Sci. 811 1-6.

x1\$

UKyl, A.C. Vyas, N.D. and Visuanath, B.	1944	Comparitive studies on Indian Soils II. Concessition of clays from soil profiles. Ind. J. Actis. Soi. 14: 345-263.
Unnitrishnan, K	1961	The effect of rainfall and clevation on the properties of laterite colls. <u>Discertation culsitted to and accepted by the University of</u> Madran.
Vandor Morwo, C	1636	Grey ferruginous laterite soils frans.3. Int. Carg. <u>Soil. Sci. 1</u> 03-306.
Vanziji, J.P.	1934	Soil problems of South Africa Soil, Res. 4: 160-167.
Vijay Ghandran, P.K.	1963	Effect of Alevation and Rainfall on forms of principal plant nutrient elements in Kerala Soils. <u>Bisseriation submittee</u> to and accorded by the University of Hadras.
Viswanath, B. Ukyl, A.C.	1944	Comparitive studies on Indian soils Regional and environmental factors associated with Indian Soils. <u>Ind. J. Aarle. Sci.</u> , <u>14</u> 333-334.
Walker, T. J. Adams, A.F.R.	1058	Studies on soil organic matter 2. Influence of increased leaching at various stages of weathering on lovels of Corbon, Hitrogen, and organic and total phosphorus. Soil. Act. 57: 1-10.

x111

Wilcox 1939 Soil moisture studies I. Some factors affecting the moisture holding capacity and its determination. Sci. Agric. 20: 140-149. 1961 Wild. A. A pedological study of phosphorus in twelve coils derived from granite. Aust. J. Agric. Res. 121 286-299. Wohltmann, P. Quoted by Das, S. in Vertical distribution of 2901 phosphate in Calcareous Soils. Ind. J. Agric. Sol., 16: 104-168. Yadav, J.S.P. Phosphorus status of Fathak 1963 cortain forest soils of India. J. Ind. boc. of soil Boi. 11: 181-187.

* Originals not seen

ΧĨΫ