

## **CATION EXCHANGE CAPACITY IN RELATION TO THE MECHANICAL COMPOSITION AND ORGANIC MATTER STATUS OF SOME SOIL PROFILES OF KERALA\***

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The influence of mechanical composition of the soil and its organic matter status on the cation exchange capacity have long been recognised. The relationship between the cation exchange capacity and the different size fractions vary considerably, increasing from coarse sand to clay. Depending on the clay content and its composition cation exchange values show wide variations, Hence the present investigation was undertaken to study how the cation exchange capacity is correlated with mechanical composition and organic matter status with regard to soils of Kerala State.

### **Materials and Methods**

Fourteen soil profiles representing the major soil groups of Kerala were collected from different parts of the State. The samples were collected in each case from fixed depths of 0 - 30 cm, 30 - 60 cm and 60 - 90 cm and in all 40 soil samples were collected.

The samples were air - dried for a week, gently ground with a wooden mallet and passed through a 2 mm sieve. The material passing through the sieve was stored in labelled bottles and used for analysis.

Mechanical analysis, organic carbon and cation exchange capacity of the soils were carried out as per analytical procedures outlined by Piper (1950).

### **Results and Discussion**

The results of analysis carried out are presented in Table 1. It may be observed that the texture of the soils studied varied from loamy sand to clay which in turn has resulted in wide variations in the cation exchange capacity. The sandy soils (Nos. 1 to 6) had a characteristically low cation exchange capacity (1.52 to 2.24 m. e/100 g) obviously due to the low content of clay present in them (7.4 to 8.6%).

The soils of the kole area (Nos. 7 to 9) have a cation exchange capacity ranging from 4.86 to 9.84 m. e/100 g. The surface horizon shows a high

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organic matter content of 4.9% compared to the lower layers. The clay content is also maximum in the surface layer and hence the high cation exchange capacity of the surface soil can be attributed to the combined effects of clay and organic matter. Turner (1935) cited by Mukherjee (1944) observed that the organic matter associated with clay had a comparatively higher cation exchange capacity. The present observation substantiates the above assumption.

Of the two alluvial profiles studied, profile IV had a comparatively higher cation exchange capacity, the surface layer showing the maximum value. The higher organic matter content of the surface layer is a major contributing factor to the cation exchange capacity of the whole soil. The decrease in organic matter content with depth showed a corresponding decrease in cation exchange capacity also. Moreover the positive correlation between clay and cation exchange capacity (+0.845) as well as between organic matter and cation exchange capacity (+0.983) bears ample testimony to the fact that clay and organic matter are the important factors affecting the cation exchange capacity of these soils. In profile number V the value of cation exchange capacity is low. The mechanical composition data show that it is dominated by the fine sand fraction which has a low cation exchange capacity. The slightly higher value observed for the surface soil is due to the organic matter content, since the amount of clay is the least in this layer. This agrees with the findings of Yuan *et al.*, (1967) who observed that in sandy soils it is the organic matter which decisively influences its cation exchange capacity. In this case since the value of cation exchange capacity is low the contribution of organic matter becomes appreciable.

The laterite soils (Nos. 16 - 24) have a low cation exchange capacity. Of the three profiles studied the range in values of cation exchange capacity is between 1.62 to 6.84 m.e/100g. In spite of the high content of clay (35 to 62.5%) the cation exchange capacity value is low indicating that it is not only the quantity but also the type of clay mineral present which plays a major role in deciding the exchange capacity. Kelley (1948) and Sathyanarayana and Thomas (1967) working on Malabar laterites, Menon and Mariakulandai (1957) working on red and laterite soils of Tamil Nadu and Koshy (1962) working on Kerala soils have obtained results which are in agreement with the present observation. The surface horizon possessed a high cation exchange capacity clearly indicating the role played by organic colloids. The present observation is also in line with that of Menon and Mariakulandai (1957). The increase in clay content with depth showed corresponding increase in cation exchange capacity values. This observation is in harmony with the findings of Roy and Landey (1962) in the red and laterite soils of Mand water-shed. The correlation coefficient worked out between clay and cation exchange capacity is positive ( $r=0.521$ ) and that between cation exchange capacity and organic matter is also positive ( $r=0.417$ ). The cation exchange capacity being generally low for these soils the contribution of the organic fraction becomes apparent. So an

**Table 1**  
**Mechanical composition and cation exchange capacity of soils**

Sl. No.	Profile No	Location & soil group	Depth (cm)	Organic matter %	Clay %	Silt %	Coarse sand %	Fine sand %	Cation exchange capacity m.e/100g
1	2	3	4	5	6	7	8	9	10
1-3	Profile I	Kayamkulam sandy	0-30	0.3	7.4	4.0	58.2	30.1	2.12
			30-60	0.2	7.7	4.8	58.5	28.8	2.24
			60-90	0.2	8.6	5.0	58.5	27.7	1.97
4-6	Profile II	Kayamkulam sandy	0-30	0.3	7.4	4.4	60.6	27.3	1.52
			30-60	0.3	7.9	4.9	60.5	26.4	1.88
			60-90	0.2	8.4	4.9	60.3	26.2	2.08
7-9	Profile III	Trichur Kole land	0-30	4.9	28.6	4.0	31.3	31.2	9.84
			30-60	0.6	28.5	5.2	32.2	33.5	4.88
			60-90	0.3	27.6	4.5	34.8	32.8	5.68
10-12	Profile IV	Alwaye Alluvial	0-30	7.2	25.7	21.0	5.8	40.3	18.84
			30-60	4.4	30.7	19.3	3.5	41.9	13.40
			60-90	3.9	31.3	19.5	2.8	42.5	13.84
13-15	Profile V	Chirayinkil Alluvial	0-30	0.3	11.3	5.8	39.4	43.2	4.40
			30-60	0.1	15.3	4.0	40.3	40.3	2.40
			60-90	0.3	15.0	5.4	38.8	40.5	2.00
16-18	Profile VI	Perinthalmanna Laterite	0-30	0.6	35.0	6.5	37.6	20.3	5.40
			30-60	0.9	61.5	6.2	22.3	9.1	6.84
			60-90	0.3	48.2	8.1	24.8	18.6	3.82
19-21	Profile VII	Angadipuram Laterite	0-30	1.6	41.1	2.5	35.1	19.7	5.10
			30-60	0.7	62.5	4.2	22.1	10.5	6.24
			60-90	0.3	60.0	4.0	20.9	14.8	5.42
22-24	Profile VIII	Vellayani Laterite	0-30	0.6	39.5	3.1	34.6	22.2	3.50
			30-60	0.3	40.4	3.3	34.3	21.7	1.84
			60-90	0.4	49.4	3.3	34.8	12.1	1.62
25-26	Profile IX	Vechoor Kari	0-30	13.9	52.4	26.5	1.4	5.8	24.32
			30-60	18.3	50.9	24.5	1.6	4.7	26.12
27-28	Profile X	Vechoor Kari	0-30	17.3	50.0	24.1	1.8	6.8	23.84
			30-00	16.6	53.4	24.5	1.0	4.5	23.36
29-31	Profile XI	Kozhinjampara Black	0-30	0.9	53.5	21.8	5.6	18.2	48.0
			30-60	0.5	57.0	19.6	5.1	17.8	47.3
			60-90	0.6	57.5	20.1	4.7	17.1	46.92
32-34	Profile XII	Kozhinjampara Black	0-30	0.4	57.0	19.2	5.6	17.8	47.16
			30-60	0.7	59.03	17.1	5.1	17.6	49.56
			60-90	0.6	60.7	16.9	4.6	17.2	46.96
35-37	Profile XIII	Palode Forest	0-30	1.8	32.5	5.1	38.9	21.7	5.9
			30-60	1.1	35.5	8.2	38.7	16.5	3.6
			60-90	0.3	37.0	8.7	39.7	14.3	3.0
38-40	Profile XIV	Patchalloor Red	0-30	0.7	27.1	2.5	40.2	29.5	2.12
			30-60	0.5	28.0	2.2	40.8	28.5	2.54
			60-90	0.3	28.5	1.9	40.8	28.5	2.4

increase in the two components will definitely increase the cation exchange capacity of the soil. This is quite clear from the results of the present investigation.

Kari soils (Nos. 25 - 28) possess a comparatively high cation exchange capacity ranging from 23.4 - 26.1 m.e./100g which agrees with the values obtained by Kurup (1967) working with similar soils. Kari soils have a characteristically high content of organic matter. In fact these soils possessed the highest organic matter content (13.9 — 18.3%) among the soils studied. Apart from this the clay fraction of these soils is said to have a preponderance of a mixture of illitic and montmorillonitic type of clays (Pillai, 1964). Hence the high cation exchange capacity of these soils can be attributed to a high content of organic matter coupled with the presence of clay mineral of the type mentioned above. The increase in cation exchange capacity of the sub surface layer is evidently due to the higher organic matter content of that layer.

The highest value of cation exchange capacity (46.92 - 49.56 m.e./100g) observed in the present investigation is in the black soils (Nos. 29 - 34). High values for cation exchange capacity of black soils have been reported by Alexander and Durairaj (1968). Ray Choudhuri *et al* (1943) observed 2:1 bidentate type of clay mineral in the clay fraction of these soils. Hence the high cation exchange capacity of these soils can be attributed to the predominance of 2:1 minerals. The more or less uniform trend in the clay and cation exchange capacity shows that very little eluviation has taken place. These soils have a comparatively low level of organic matter and hence contribution of organic colloids in the cation exchange reaction is only of minor importance. Decrease in cation exchange capacity with depth is noticed in profile XI, while increase in intermediate layers is noticed in profile XII. Das *et al.* (1946) working with Indian soils also made similar observations with regard to variations in cation exchange capacity down the profile.

The forest soils (Nos. 35 - 37) show a very low value of cation exchange capacity (3 - 5.9). The deeper layers tend to be laterite with characteristically of the surface and sub surface layers appears to be the contributing factor in the high cation exchange capacity of these layers. The clay content shows an increasing trend downward with no corresponding change in cation exchange capacity, clearly indicating that the clay fraction has low cation exchange capacity.

Red soils in general have a low cation exchange capacity. The profile studied in the present investigation also conforms to this. The results obtained agree with the findings of Lall (1955) in the red soils of Bihar and Menon



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